

1982

STRATIGRAPHY, STRUCTURE, AND ECONOMIC GEOLOGY

Introduction

This geological investigation of the Sheridan quadrangle was conducted as part of a regional stratigraphic study in the northern Oregon Coast Range to assess the potential for oil and gas.

The Sheridan quadrangle was first investigated by Baldwin and others (1955) as part of the U.S. Geological Survey oil and gas studies. Baldwin and Roberts (1952), Baldwin (1964), MacLeod (1969), Brownfield and Schlicker (1981a,b), and Brownfield (1982a,b) mapped areas adjoining the Sheridan quadrangle on the east, north, south, and west. McWilliams (1968, 1970) included the Sheridan area in his regional study of the Paleogene stratigraphy and tectonics of the Oregon Coast Range.

Stratigraphy

The oldest rocks exposed in the Sheridan quadrangle are the Siletz River Volcanics, a thick sequence of basalt flows, pillow basalt, flow breccia, and massive to finely bedded basaltic fragmental debris intercalated with volcanoclastic marine sedimentary rocks of early to middle Eocene age. The unit was first described by Shively and Baldwin (1949) and later redefined by Shively, MacLeod, and Wagner (1965). It is the oldest exposed unit in the Coast Range, and nowhere in the central or northern portions is its base exposed.

No fossils have been collected in the Siletz River Volcanics in the Sheridan quadrangle. Molluscan faunas collected by Baldwin (1964) and MacLeod (1969) south and west of the mapped area were assigned to the Cagay Stage of Weaver and others (1944). Several foraminiferal assemblages were collected in the Siletz River Volcanics outside of the quadrangle by MacLeod (1969) and McWilliams (1968, 1973) and were assigned to the Ulianian and Penian Stages of Mallory (1959).

Regionally, the upper surface of the Siletz River Volcanics is irregular because of volcanism, tectonism, and erosion. The volcanic pile influenced the depositional patterns of disconformably overlying middle to upper Eocene sediments, whose ages vary depending upon their relative positions along the flanks of the volcanic pile.

Marine sediments consisting of carbonaceous, micaceous shale and siltstone with local interbeds of gray to greenish-gray basaltic sandstone of the Yamhill Formation disconformably overlie the Siletz River Volcanics. The Yamhill Formation was first described by Baldwin and others (1955); its type locality is along Mill Creek from a point just south of the Sheridan quadrangle north to the Mill Creek School.

Megafossils are rare in the Yamhill, but most of the unit contains abundant Foraminifera which indicate a middle and upper Eocene age. Several foraminiferal assemblages were collected by Baldwin and others (1955) and McWilliams (1968, 1973) within the Sheridan quadrangle. Fossil assemblages collected by the author were identified and assigned by Michael (1982, written communication to the Narrian Stage of Mallory (1959). To the west of the mapped area, MacLeod (1969) collected several foraminiferal assemblages which were assigned to the lower Narrian and upper Ulianian Stages of Mallory (1959).

Tuffaceous siltstone and shale with intercalated basalt flows, pillow basalt, pillow breccia, and tuff of the upper Eocene Nestucca Formation (Shively and Vokes, 1949) unconformably overlie the Yamhill Formation. This middle upper Eocene unconformity represents a major period of deformation and erosion in the central Oregon Coast Range.

Several foraminiferal assemblages were collected in the Nestucca Formation within the Sheridan quadrangle (Baldwin and others, 1955; McWilliams, 1968, 1973) and to the west of the study area. MacLeod (1969) and assigned to the upper Narrian Stage (Mallory, 1959). Along the Oregon coast, the foraminiferal assemblages within the Nestucca Formation are indicative of the lower Redanian (Schneek and Klempf, 1936; Klempf, 1938) and upper Narrian Stages (Shively and others, 1965).

Structure

The central part of the Oregon Coast Range has a northward-trending anticlinal form; the Sheridan quadrangle is located on its east flank. A broad east-west downward crosses the Coast Range just north of latitude 45° N. and forms a broad east-west plunging fold transverse to the Coast Range antiform. Small amplitude folds are common along streams where the sedimentary rocks are well exposed, but due to limited exposures they cannot be traced beyond the stream exposures.

The Yamhill River fault of Baldwin and others (1955), a major northeast-trending fault, roughly parallels the South Yamhill River valley and juxtaposes the Siletz River Volcanics, Yamhill Formation, and Nestucca Formation. This fault, which apparently is down on the north side, is a continuation of a fault first mapped by Baldwin and Roberts (1952) in the southern part of the Grand Ronde-Spirit Mountain quadrangle. MacLeod (1969) indicates the fault extends at least 15 mi west of the Sheridan quadrangle. A late Eocene age for the fault is consistent with the structural relationships in the Grand Ronde quadrangle, where late Eocene diabase is intruded along the Yamhill River fault. Regional studies indicate the possibility of right-lateral motion along this fault. Two other northeast-trending faults in the southwestern part of the Sheridan quadrangle may be related to stress adjustments between the Yamhill River fault and a second major east-west trending fault just south of the quadrangle boundary. Numerous, apparently small faults are seen cutting sedimentary rocks along streams, and large changes in attitudes of strata may relate to nearby faults now buried by alluvium or not identifiable in the poor exposures typical of much of the area.

Economic Geology

The Yamhill Formation contains abundant organic matter and could be considered a possible source rock for petroleum. Most of the sandstone interbeds have low porosity resulting from the presence of clay altered volcanoclastic material. These beds would not be satisfactory reservoir rocks. The overlying tuffaceous Nestucca Formation lacks the necessary qualities for both petroleum generation and accumulation. No oil and gas holes have been drilled within the quadrangle to test oil and gas potential. Stratigraphic relationships evident in the quadrangle, however, should aid regional hydrocarbon assessments.

Quarry rock is abundant within the Sheridan quadrangle. Intrusive rocks, intercalated basalts in the Nestucca Formation, and the Siletz River Volcanics are good sources for crushed rock. Gravels that occur near the base of the river terrace deposits have been used for road construction.

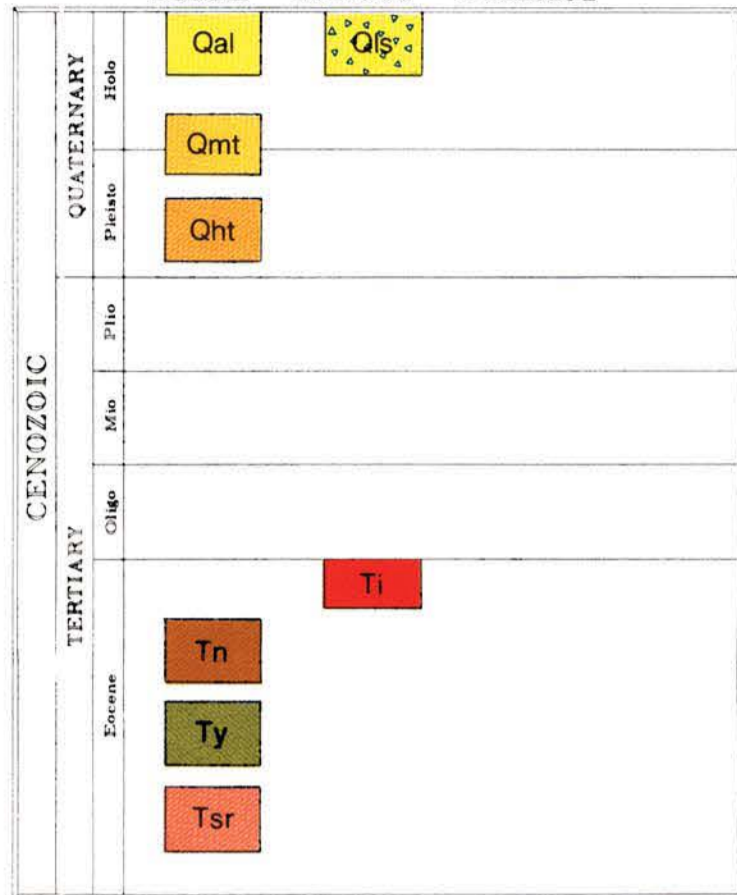
There are known clay deposits in the quadrangle. One deposit, at least 50 ft thick (Wilcox, 1935), was mined near the town of Yamhill by the Williams Clay Products Company until 1976. The unfired clay in this deposit varies in color from light gray to black. The black-colored clay contains a large amount of carbon which, when oxidized during firing, produces a nearly white brick.

The clays are developed within the Yamhill Formation on the erosional surface below the unconformity with the overlying Nestucca Formation. Similar deposits may be found elsewhere in the area along this contact.

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TIME ROCK CHART



EXPLANATION

SURFICIAL DEPOSITS (HOLOCENE AND PLEISTOCENE)

Age ranges of individual units overlap

- Qal** Alluvium and lower terrace deposits—Mostly poorly sorted, unconsolidated deposits of clay, silt, sand, and fine to very coarse gravel; includes recent alluvial terrace deposits along the South Yamhill River and its tributaries. 5-80 ft thick.
- Qht** High terrace deposits—Weathered, poorly sorted, semi-consolidated deposits of gray to reddish-brown clay, silt, sand, and fine to very coarse gravel; massive to finely bedded and tilted. Wilamette Silt by Baldwin and others (1955). 0-100 ft thick.
- Qmt** Middle terrace deposits—Poorly sorted, semi-consolidated deposits of clay, silt, sand, and fine to very coarse gravel; may include thin deposits of light brown, massive to finely bedded and tilted Wilamette Silt by Baldwin and others (1955). 0-100 ft thick.
- Qls** Landslide deposits—Unconsolidated, poorly sorted slide deposits with hummocky surfaces. Mapped only where readily apparent; small soil and rock failures not shown. Nestucca (Tn) and Yamhill (Ty) strata are particularly susceptible to landsliding.

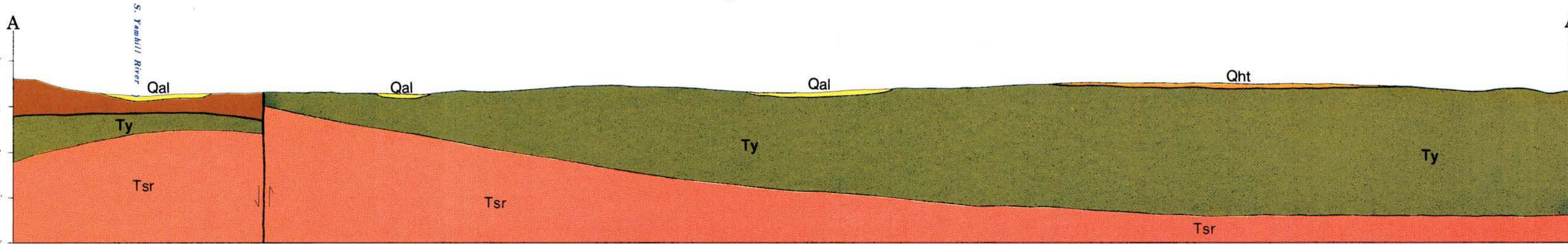
BEDROCK GEOLOGIC UNITS

- Ti** Diabase intrusive rocks (upper Eocene)—Basalt and diabase dike and sill. Medium to coarse-grained hypocrystalline to holocrystalline basalt and diabase, with subophitic to ophitic, porphyritic, or less commonly, intergranular and interstitial texture. Consists of 40-55 percent labradorite, 10-35 percent augite, and minor amounts of olivine, magnetite, apatite, and zircon.
- Tn** Nestucca Formation (upper Eocene)—Light brown to light gray tuffaceous siltstone and shale, with minor light gray tuffaceous, ridgipathic, cross-bedded sandstone near the base. Contains interbedded basalt flows, pillow basalt, pillow breccia, and tuff. Perforated by basalt dikes (see locally present). Unconformable above the Yamhill Formation. Foraminiferal faunas were assigned to the upper Narrian Stage of Mallory (1959) by McWilliams (1968, 1973), and molluscan faunas were referred to the upper Eocene (Tyne Stage) by Baldwin and others (1955) and McWilliams (1968). At least 3,000 ft thick.
- Ty** Yamhill Formation (middle and upper Eocene)—Medium to dark gray, massive to finely bedded, micaceous, carbonaceous, tuffaceous shale and siltstone, with locally interbedded medium gray to greenish-gray calcareous, fossiliferous, basaltic sandstone, minor limestone concretions. Foraminiferal assemblages were assigned to the Narrian Stage of Mallory (1959) by McWilliams (1968, 1973) and McKee (1982, written communication), and molluscan faunas were assigned to the late Eocene by Baldwin and others (1955). As much as 3,000 ft thick.
- Tsr** Siletz River Volcanics (lower and middle Eocene)—Dark greenish-gray, aphanitic to porphyritic, vesicular basalt flows and pillow basalt, basalt breccia, and tuffaceous, with interbedded red to green calcareous sandy tuff; contains medium to dark gray, calcareous, tuffaceous shale, siltstone, and sandstone at top of section. Foraminiferal assemblages were assigned to the Ulianian and Penian Stages of Mallory (1959) by McWilliams (1968, 1973), MacLeod (1969), and McKee (1980). Megafossils were assigned to the Cagay Stage of Weaver and others (1944) by Baldwin (1964) and MacLeod (1969). Maximum exposed thickness in the quadrangle is approximately 500 ft. Total thickness unknown but probably exceeds 12,000 ft locally in the Central Coast Range.

GEOLOGIC SYMBOLS

- Contact—Approximately located and inferred; contacts exposed only along streams and roads.
- - - Fault—Approximately located; dashed where inferred; dotted where concealed. Bar and ball on downthrown side.
- Strike and dip of beds—Attitudes along stream beds were measured on bed rock that is intermittently exposed or too small to show at the scale of the map. Units are shown in parentheses north of the Yamhill River fault. Yamhill Formation is exposed in Mill Creek south of fault.
- Included Horizontal
- Fossil locality—References and age given in the table.

Geologic Cross Section



Geology by Michael E. Brownfield; some bedrock attitudes after Baldwin and others, 1955

Field work completed in 1981

Base Map by U. S. Geological Survey

Control by USGS, USC&GS, USCE, and State of Oregon

Topography from aerial photographs by photogrammetric methods

Aerial photographs taken 1954. Field check 1956

Polynomial projection, 1927 North American datum

10,000-foot grid based on Oregon coordinate system, north zone

1000-meter Universal Transverse Mercator grid ticks, zone 10, shown

Also prepared by

STATE OF OREGON

DEPARTMENT OF GEOLOGY AND MINERAL INDUSTRIES

Cartography by Paul E. Staub, 1982

SCALE 1:24,000

1000 0 1000 2000 3000 4000 5000 6000 7000 FEET

1 0 1 2 3 4 5 6 7 8 9 10 KILOMETER

CONTOUR INTERVAL 20 FEET

DOTTED LINES REPRESENT 10-FOOT CONTOURS

NATIONAL GEOLOGIC VERTICAL DATUM OF 1929

UTM GRID AND 1970 MAGNETIC NORTH DECLINATION AT CENTER OF SHEET

OREGON

QUADRANGLE LOCATION