

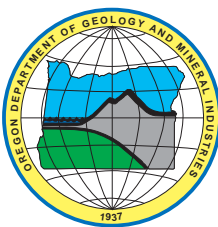
State of Oregon
Department of Geology and Mineral Industries
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Open-File Report O-07-09

**PRELIMINARY GEOLOGIC MAP OF THE BROWN MOUNTAIN 7.5' QUADRANGLE,
JACKSON AND KLAMATH COUNTIES, OREGON**

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2007

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Oregon Department of Geology and Mineral Industries Open-File Report O-07-09
Published in conformance with ORS 516.030

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Introduction

The Brown Mountain Quadrangle (BMQ) is located directly on the north - south oriented axis of High Cascade volcanoes that stretch from British Columbia, Canada to Mount Lassen in northern California. Figure 1 provides exact location information on several levels; where in the Pacific Northwest to where in south-central Oregon and to the local geographic context in which the BMQ is situated. Topographically the elevation ranges in the BMQ from 4190 feet where the South Fork of the Little Butte Creek exits the quadrangle flowing to the west, to 7311 feet, the summit of Brown Mountain, the highest point in the quadrangle. That means there are 3121 feet of topographic relief within the BMQ. Naturally anyone who has either read about or visited regions within the higher portions of the Cascade Mountains knows that the landscape is an interesting mix of volcanic and glacial features, often referred to by the imagination-inciting phrase “fire and ice”. However, in the BMQ physical and visual evidence of the “ice” part of the phrase is relegated to the very highest segment of Brown Mountain, particularly on the north and northeast facing slopes. The “fire” part of the phrase is most easily visualized in the northeastern quarter of the quadrangle that is dominated by numerous blocky andesite flows that radiate outward from the summit. As andesite stratovolcanoes go Brown Mountain is quite atypical because the proportionality of lava flows to pyroclastic material is highly skewed towards lava flows.

All the other topographic high points in the BMQ are less than 6200 feet in elevation. Prominent volcanic vents like Robinson Butte, Burton Butte, and Old Baldy (see Figure 1) extruded much lava over large areas of the BMQ. These high points have one thing in common. They are all located just outside of the quadrangle boundaries. Cox Butte and Brush Mountain, which extruded basaltic andesite and basalt respectively, are two prominent named volcanic vents that reside within the quadrangle. Several cinder pits / rock quarries used exclusively for road building aggregate are found scattered through the quadrangle, the largest of which is the Big Elk cinder pit located in the extreme northwest portion the quadrangle, approximately one mile SSE of the Robinson Butte summit.

The Pacific Crest Trail crosses the BMQ from south to north (see Figure 1). There are several points of easy access to the PCT. One is where Oregon highway 140 crosses the PCT north of Brown Mountain and south of Mt. McLoughlin and another is where the Dead Indian Memorial highway crosses the PCT. At both locations small paved parking lots are available for vehicles. Also the trail from Fish Lake to Lake of the Woods intersects the PCT which provides through hikers an opportunity to re-supply and use the shower and laundry facilities at the small stores located at both popular fishing locations.

A comment on nomenclature: when geoscientists classify igneous rock samples they often come at it from two points of view. One is based on identifying the visible minerals in a hand sample (a modal mineral classification) and the other is based on a chemical analysis of that sample (a chemical classification of igneous rocks – see Figure 2 as an example). Naturally the latter is more precise and rigorous and the former is looser and less precise and is open to more opinions. The most common volcanic rock names (basalt, basaltic andesite, andesite, dacite, rhyolite) define a sequence in which the iron – magnesium

bearing silicate minerals (olivine, orthopyroxene, clinopyroxene, hornblende, biotite) are most abundant on the left side of the sequence, forming upwards of 50 to 60 percent of the minerals present and decreases to nearly zero to the right, namely, in rhyolite. The remaining 40 to 50 percent of the rock consists mostly of plagioclase feldspar, non-iron magnesium bearing silicate mineral, and a few percent of chromium, iron, titanium dominated oxide minerals. With regard to rock chemistry silica (SiO_2) increases from basalt to rhyolite and correlates directly with increasing viscosity and greater explosivity.

Table 1, which accompanies the geologic map of the BMQ, contains the chemical and age data for all the analyzed rock samples. Figure 1 also depicts the location of all the samples for which age dates exist, both within the BMQ as well as immediately adjacent to it. Figure 2 is a total alkali ($\text{Na}_2\text{O} + \text{K}_2\text{O}$) versus SiO_2 diagram that summarizes the rock names that are most germane for the volcanic materials present in the BMQ. In addition the chemical data is displayed for each stratigraphic unit that is defined below using an individualized symbol that is summarized in the legend that accompanies Figure 2.

Broad Overview of the Geology of the Brown Mountain Quadrangle

The geology of the BMQ encompasses the latest 22 million years (Ma) of earth history. To be scrupulously honest it is more accurate to say 99+ percent of the rocks outcropping within the BMQ are 6.2 Ma old or younger with just a small sliver of the early Miocene Heppsie Formation. These Miocene-age volcanic rocks are exposed in the South Fork of the Little Butte Creek drainage just before the stream crosses into the adjoining Robinson Butte quadrangle on the far western margin of the BMQ. The Heppsie Formation is the youngest segment of the Little Butte Volcanic Series that in turn is classified as part of the Late Early Western Cascade period of volcanism as defined by Priest (1990). The uppermost section of the Heppsie Formation in this region consists of a poorly to moderately welded ash-flow tuff unit that is well exposed in the north valley wall of the South Fork of the Little Butte Creek in the eastern one-third of the Robinson Butte quadrangle. Locally this ash flow unit is several hundred feet thick and takes on a crude columnar jointing pattern where welding has been moderate. To accurately picture how this material was both erupted and deposited on the earth's surface, one needs to envision the famous eruption of Vesuvius volcano in Italy in 79 AD as described by Pliny the younger, the eruption and near total destruction of Mount Mazama, nearly 7000 years ago, or the much more recent eruption of Mount Pinatubo in the Philippine Islands in 1991.

Resting on top of the Heppsie Formation in this region with an unconformity or age gap of 13 to 15 million years are lavas that belong to the early segment of High Cascades volcanism as summarized by Priest (1990). In this portion of the Cascades Mountains latest Miocene to Holocene volcanism is completely dominated by lavas and vent pyroclastics. Whereas pyroclastic flow activity characterized the youngest segment of Heppsie time, the latest 6 to 7 Ma of geologic time has seen very little of this kind of violent volcanic activity with the one very notable exception being Crater Lake / Mount Mazama volcano and its pyroclastic flow dominated eruptive history ~7000 years ago. Geologically-speaking the extrusive rocks that are so ubiquitous in the quadrangle are latest Miocene (7.246 to 5.332 Ma), Pliocene (5.332 to 1.806 Ma), and Pleistocene (1.806 to 0.0115 Ma) in age (see Gradstein, Ogg, and Smith, 2004 for details). Mertzman (2000) and Mertzman (unpublished

data, 2007) provide many new age dates, derived from both a whole rock K-Ar method as well as $^{40}\text{Ar}/^{39}\text{Ar}$ technology, that have been measured through June, 2007.

Glacial activity over the past 1 to 2 Ma has only affected the very highest point in the quadrangle; namely, the summit region of Brown Mountain volcano. Evidence for glacial activity can be seen on the north and northeast segment of the Brown Mountain summit crater.

Explanation of Map Units

Qal Alluvium (Holocene) Unconsolidated sediment found in close proximity to modern drainages.

Volcanic Rocks

Qbv, Qav Basaltic to basaltic andesite / andesite vent deposits (Pleistocene) Poorly lithified to unconsolidated lapilli to ash-sized cinders black to brown to red in color with lesser amounts of similarly colored lava spatter, bombs, and scoria. These deposits mark volcanic vents areas that are often cinder cones.

Qabm Andesite of Brown Mountain (Upper Pleistocene) Numerous blocky lava flows of aphanitic andesite originate from Brown Mountain. The outermost carapace of the blocky flow material is medium to dark gray in color and dominated by open vesicles that are 1 to 2mm in diameter and can be stretched out into an ellipsoid in the flow direction. The flow interiors where visible are significantly lighter in color and are universally characterized by platy flow jointing with 2 to 3 cm separating the joint planes. Hand samples contain 10 to 20 percent microphenocrysts between 0.4 and 1 mm, with plagioclase feldspar being the most abundant mineral, followed by orthopyroxene and clinopyroxene. Minor olivine and opaque oxide minerals are also present. The age of the Brown Mountain volcanic activity is less than 100,000 years old, most likely between 50,000 and 25,000 years old based on geomorphologic evidence. Vestiges of Late Pleistocene glacial and periglacial activity are confined to the summit region of Brown Mountain, in particular, on the northern and northeastern sides just below the summit crater.

Qbbb Basalt of Burton Butte (Middle Pleistocene) Light gray to dark bluish gray in hand sample color, lava samples are consistently lighter in color than that of pyroclastic samples. Burton Butte cinder / scoria cone, the source of these lavas, is located in the adjoining quadrangle to the east, the Lake of the Woods, South quadrangle. Pahoehoe lava flows from Burton Butte spread westward nearly six miles, down the paleo-drainage now occupied by the Beaver Dam Creek all the way to Deadwood Prairie. Most samples have a diktytaxitic (sponge-like) texture with a larger sized set of vesicles present, several mm to one centimeter in diameter, that are often lined to partially filled with secondary mineralization, mostly carbonate with some silica and zeolitic minerals infrequently present. The larger sized set of vesicles is often stretched out to provide a lineation parallel to the last flow direction

of the lava. Plagioclase, 0.5 to 2mm in diameter, is the most abundant mineral with 15 to 25 percent olivine and a similar amount of clinopyroxene present, filling in the interstices between the tabular laths of plagioclase. Chromite, present within early-formed olivine crystals, together with titanomagnetite and ilmenite, are the opaque minerals that constitute 8 to 10 percent of the minerals present in these basaltic lavas. Several whole rock K-Ar ages are available for this unit and are characterized by relatively large + or – values. These relatively large error limits are due in part to the low whole rock K₂O values for the Burton Butte basalt samples, <0.3%. One ⁴⁰Ar/³⁹Ar age has been quite recently determined and is preferable over all the others. Burton Butte volcanic activity is 0.33 ± 0.12 Ma old.

Qbrb Basalt of Robinson Butte (Middle Pleistocene) Light gray to dark bluish gray in hand sample color, lava samples are consistently lighter in color than that of pyroclastic samples. Robinson Butte cinder / scoria cone, the source of these lavas, is located in the adjoining quadrangle to the west, the Robinson Butte quadrangle. Three to five meter thick lava flows with pahoehoe surfaces are abundant. Vesicles 5 to 10 mm long and stretched out in the direction of flow and partially lined with vapor phase deposited secondary minerals, are common. The lavas have a glomeroporphyritic texture with olivine (1 to 3 mm in maximum dimension) constituting 10 to 15 percent of each hand sample. Poikilitic within early-formed olivine crystals are chrome-spinel inclusions. Abundant clinopyroxene (1 to 3 mm) and plagioclase (0.5 to 1 mm) phenocrysts constitute nearly 50 percent of a hand sample and are present in clumps as well as individual crystals. The groundmass is primarily intersertal in nature. One whole rock K-Ar age is available for this unit, 0.40 ± 0.30 Ma. One ⁴⁰Ar/³⁹Ar age has been quite recently determined and is preferable because of its much better precision. Robinson Butte volcanic activity is 0.36 ± 0.06 Ma old. From a geologist's or hiker's point of view it is very interesting to note that the volcanic activity at Robinson Butte and Burton Butte are virtually identical in age, which suggests these two volcanoes could have simultaneously been erupting basaltic lava onto the southern Oregon landscape. Since the chemical compositions of basaltic lavas from Robinson and Burton Buttes are easily distinguishable, one implication is that the plumbing systems for each of these two volcanoes are separate and distinct.

Tpbm Basalt of Brush Mountain (Upper Pliocene) The Brush Mountain volcanic vents are aligned in a NNW-SSE direction in the extreme south-central part of the BMQ, extending into the Little Chinquapin Mountain quadrangle. Numerous basaltic lava flows are found in both quadrangles that emanate from these linearly arrayed vents. Brush Mountain basalts are medium gray in color as lavas and considerably darker gray as the lavas become more vesicular in nature. These extrusive rocks are porphyritic with 10 to 15 percent phenocrysts with plagioclase feldspar somewhat more abundant than olivine. These two minerals occur separately and in glomeroporphyritic clumps that can range up to 3 mm in diameter. Pyroxene is mostly confined to the matrix. The olivine is partially altered to iddingsite, a characteristic that causes the olivine phenocrysts to be much more darkly colored than the typical lime green color; it also causes olivine to be iridescent, particularly in shades of purple on newly broken surfaces. Two whole rock K-Ar ages are available for this unit, 2.22 ± 0.12 Ma and 1.97 ± 0.08 Ma old.

Tpbao Basaltic Andesite of Old Baldy (Middle Pliocene) The singular topographic high point known as Old Baldy is located at the western tip of an east-west oriented ridge that on its eastern extent is composed of latest Miocene hornblende andesite. Earlier in the twentieth century a USFS fire lookout tower was located at the top of Old Baldy. Presently the PCT crosses this ridge several hundred meters to the west of Old Baldy. This ridge is the drainage divide between the South Fork of the Little Butte Creek and its tributaries which eventually end up in the Rogue River and the streams that drain Johnson Prairie to the south that empty into Jenny Creek and then the Klamath River. Every hand sample of Old Baldy basaltic andesite has 5 to 10 glomeroporphyritic clumps, 5 to 8 mm in diameter, that are composed of plagioclase, far and away the most abundant mineral present in the Old Baldy lavas, and olivine. The abundance of olivine diminishes rapidly once the microphenocrysts and the matrix are examined because pyroxene, particularly clinopyroxene, becomes the dominant mafic mineral present. Acicular to lath-like plagioclase in the matrix of the samples is a particularly noticeable trait when examining them with a hand lens. Two whole rock K-Ar ages are available for this unit, 2.78 ± 0.10 Ma and 2.63 ± 0.32 Ma old.

Tpbcb Basalt of Cox Butte (Middle Pliocene) The vents of the Cox Butte volcano are located somewhat west of center in the BMQ. It abuts the voluminous basaltic andesite to andesite lava flows of the somewhat older Ichabod Spring Formation to the north. Cox Butte lavas are light to medium gray in color, often with conspicuous flow jointing present with 5 to 10 cm separating each of the shear planes. Three to five percent clinopyroxene and olivine phenocrysts are visible, each up to 3 mm long, with the olivine being partially to completely altered to iddingsite. Tabular shaped crystals are lined up to produce a flow lineation. Plagioclase, which is most abundant mineral present, is ≤ 1 mm in diameter, effectively concentrated in the matrix. One very interesting petrographic feature concerning Cox Butte lavas concerns the larger clinopyroxene crystals. When examined with the petrographic microscope rather than being an individual crystal, the clinopyroxene grains are mosaics of dozens of smaller pyroxene crystals, resembling bathroom tile as a first-order comparison. These multi-grained crystals may not actually be Cox Butte magmatic material, but rather fragments of pre-existing rock that became included in Cox Butte magma on its way to the surface. One whole rock K-Ar age is available for this unit, 2.96 ± 0.06 Ma old.

Tpbdp Basalt of Daley Prairie (Middle Pliocene) The vents for the Daley Prairie lavas are a series of spatter and scoria cones aligned in a predominantly east-west orientation and are located to the east of Cox Butte. A hand-sample of Daley Prairie basalt is light to medium gray in color and has 12 to 15 percent green olivine phenocrysts that are 1 to 3 mm in diameter. The olivine crystals occur singularly and in clumps creating a glomeroporphyritic texture. Some of the olivine crystals are dark brown around the margins as a result of alteration to iddingsite, a secondary assemblage of minerals that form at warm but low temperature when the lava is completely solidified. Iron-rich oxides and clay minerals are the dominant constituents in iddingsite. The Daley Prairie basalt is termed an olivine-phyric basalt because all the other silicate and oxide minerals present in the extrusive rock are confined to the fine grained matrix of the sample; thus, leaving olivine as the only visible mineral in the hand specimen. In thin section the olivine crystals have

included within them small <0.1 mm in diameter crystals of spinel, a chemically complex oxide mineral that has a general formula of AB_2O_4 where A is Fe^{+2} and Mg^{+2} and B is Cr^{+3} and Al^{+3} and Fe^{+3} in various amounts. The texture indicates the spinel mineral grains crystallized from the molten magma before the olivine crystals, which eventually grew around and enclosed the spinel crystals. An early-formed mineral subsequently surrounded by a later-formed mineral forms what is termed a poikilitic texture. Olivine-rich basaltic lavas like those of the Daley Prairie unit are thought to have likely traveled to the Earth's surface from the asthenosphere where they were formed. They seemed to have been able to make this nearly 100 km trip with little in the way of chemical interaction with the rocks in which they were in continuous contact. Two whole rock K-Ar ages are available for this unit, 3.16 ± 0.12 Ma and 2.77 ± 0.05 Ma old.

Tpais Andesite of Ichabod Spring (Middle Pliocene) Whereas the physical appearance of the preceding unit is exclusively dominated by olivine phenocrysts, the Andesite of Ichabod Spring is dominated by abundant phenocrysts of plagioclase and pyroxene. A minority of lavas included within this unit is basaltic andesite in bulk composition and can often be distinguished by the presence of 1 to 2 percent olivine phenocrysts in addition to the primary constituents plagioclase and pyroxene. Twenty to thirty percent plagioclase phenocrysts, more often in lath-like elongate crystals ranging in size up to 5 mm in length with striations present parallel to the long direction of the grains, dominate the physical appearance of these lavas. Also present are dark green to nearly black clinopyroxene and dark brown to nearly black orthopyroxene phenocrysts that range in size up to 4 mm, constituting between 5 and 10 percent of the volume of a hand sample. With some frequency these larger phenocrysts have their long axes aligned parallel to the direction in which the lava was flowing producing what is termed a trachytic texture. Many of these lavas flowed from east to west across the BMQ from vents located on the western margin of the Lake of the Woods South quadrangle. Two whole rock K-Ar ages are available for this unit, 3.42 ± 0.06 Ma and 3.25 ± 0.05 Ma old.

Tpafi Andesite of Fish Lake (Lower Pliocene to Lower Pleistocene) This unit primarily consists of andesite lava flows but intercalated amongst the andesite flows are more mafic flows of basaltic andesite, much like the Andesite of Ichabod Spring. These andesites are quite porphyritic with plagioclase phenocrysts, singularly and in glomeroporphyritic clumps up to 7 and 8 mm in diameter, that amount to between 30 and 40 percent of a typical hand specimen. Both orthopyroxene and clinopyroxene are present usually amounting to 5 to 7 percent of the hand sample. Olivine is either absent or present at a level of no more than 1 to 2 percent. On the whole outcrop of this unit is relatively limited and leaves the impression that it is more weathered due to iron staining in and along fracture surfaces that pass through the rock. No vents for these lavas were discovered. Three whole rock K-Ar ages are available for this unit, 3.97 ± 0.08 Ma, 3.56 ± 0.06 Ma, and 1.67 ± 0.03 Ma old.

Tpbmp Basalt of Moon Prairie (Lower Pliocene) These olivine-phyric lavas form the far southwest corner of the map area and have flowed from vents located further to the west in the Robinson Butte quadrangle. Spheroidal weathering is encountered at most of the outcrops of this unit. Olivine forms 15 to 20 percent of a hand sample

as phenocrysts 1 to 4 mm in diameter, many of which are partially to completely altered to iddingsite. Numerous small grains of spinel are poikilitically enclosed within the olivine crystals. Plagioclase feldspar is the most abundant mineral present but is confined to the matrix as elongate rectangular (laths) that are flow aligned to produce a trachytic texture. Present as microphenocrysts and small-sized crystals clinopyroxene is the only readily identifiable pyroxene present in these light gray colored basaltic lavas. One $^{40}\text{Ar}/^{39}\text{Ar}$ age is available for this unit, 4.51 ± 0.01 Ma old.

Tmaob Andesite of Old Baldy East (Upper Miocene) This unit is atypical for this particular segment of the Cascade volcanic province in that the silicate mineral hornblende, a double chain hydrous ferromagnesian mineral is present as scattered 3 to 5 mm long phenocrysts that constitute 2 to 3 percent of a hand sample. Strong chemical zoning is apparent in plagioclase feldspar, the most abundant present in this andesite unit. It is present as singular phenocrysts and in glomeroporphyritic clumps 1 to 3 mm in diameter that constitute 10 to 12 percent of the lava and thoroughly dominates the matrix in terms of abundance. Orthopyroxene is a significant mineral as well, constituting 7 to 10 percent of the total volume of the rock in crystals ranging up to 2 mm in diameter. Once again spheroidal weathering is a ubiquitous physical feature observed at most outcrops of this unit. One whole rock K-Ar age is available for this unit, 5.77 ± 0.09 Ma old.

Tmbab Basaltic Andesite of Beaver Dam Creek (Upper Miocene to Lower Pliocene) Once again spheroidal weathering is the order of the day when it comes to examining outcrops of this unit. Nearly every outcrop has one or more very good examples of this weathering phenomenon. Individual rounded boulders will frequently display distinctive weathering rinds when broken open, with the most intensively weathered material near the outside surface of the sample becoming less intensively weathered as one penetrates further into the rock mass. Compositionally speaking this unit has a majority of lavas that are basaltic andesite in composition; however, there are minor but persistent basaltic lavas present as well. These light gray basalts are consistently of the olivine-phyric type with the flow aligned phenocrysts of olivine ranging from 2 to 4 mm in length and constituting nearly 20 percent of the volume of a hand sample. Many of these olivine crystals have been partially to substantially altered to iddingsite. Compared to these basaltic lavas the flows of basaltic andesite are much darker bluish-gray in color even though they are higher in silica content than the basaltic lavas. The basaltic andesite lavas have fewer phenocrysts. Plagioclase feldspar and olivine phenocrysts, 1 to 3 mm in diameter, constitute 3 to 5 percent of the volume of a hand sample and are nearly equal in abundance. Pyroxene and much additional plagioclase are confined to the sample matrix that is much finer grained than in the basaltic lavas. Hence, when breaking the basaltic andesite lava in to smaller chunks conchoidal fracture surfaces are very abundant. Two whole rock K-Ar ages are available for this unit, 5.82 ± 0.17 Ma and 4.60 ± 0.34 Ma old.

Tmbpb Basalt of Pole Bridge Creek (Upper Miocene to Lower Pliocene) The Basalt of Pole Bridge Creek outcrops rather spectacularly in the region near the

confluence of Beaver Dam Creek, Pole Bridge Creek, and the South Fork of the Little Butte Creek in the extreme west-central portion of the quadrangle. On the east side of the intersection the bluffs are held up by basalt lava flows and are capped by younger basaltic andesite lavas. Individual bluffs have relief ranging from 35 to 60 meters. The maximum thickness of the Basalt of Pole Bridge Creek is on the order of 70 to 80 meters. It thins rapidly to the west so that in the Short Creek drainage in the Robinson Butte quadrangle, 2 to 2.5 km further to the west, the Basalt of Pole Bridge Creek has thinned to zero and the 0.4 Ma old Robinson Butte Basalt lies directly upon early Miocene volcanic rocks of the Heppsie Formation. On the outcrop, lava flows range from platy jointed near the base of the flow, with individual plates 1 to 3 cm thick, to essentially massive in the interiors. Individual hand samples are light gray in color and often possess two generations of vesicles. The first generation of vesicles is larger, forming spherical to oval shaped voids that are 0.2 to 0.5 cm in effective diameter. The second generation of vesicles is pinhead in size and spherical in shape. Both vesicle types are quite fresh; that is, both are unfilled and unlined by secondary minerals in most locations. The flows range from nearly aphyric; that is, phenocryst-free, to containing 20 to 25 percent small phenocrysts ranging between 1 and 2 mm in maximum dimension and consisting in order of decreasing abundance plagioclase, clinopyroxene, olivine, and opaque minerals. In many instances the olivine phenocrysts have been substantially altered through high temperature oxidation and have an appearance in thin section analogous to wormy wood. The opaque minerals are unusually abundant in this formation (10 to 12 percent of total mineralogy) and crystallized early as evidenced by their inclusion within all the silicate minerals listed above. Also unusual is both the size and abundance of clinopyroxene crystals, in this case nearly equal in abundance to that of olivine. Many basalts of similar silica content from this region of the Cascades are dominated by olivine with clinopyroxene relegated to a microphenocryst and / or matrix constituent. Several whole rock K-Ar ages have been determined on lava samples from this formation. The radiometric ages range from 6.14 ± 0.15 to 4.35 ± 0.04 Ma old.

Tmvmhf Heppsie Formation (Lower Miocene) The Heppsie Formation constitutes the youngest portion of the Little Butte Volcanic Series. In this segment of the Formation's outcrop area the older extrusive rocks are dominated by basalt, basaltic andesite, and andesite lava flows whereas in the younger, upper portion, poorly to moderately welded pyroclastic flow and air-fall tuffaceous material dominate. These latter extrusive materials are andesite to dacite in terms of their bulk composition. On the outcrop the lavas are greenish-brown in surface coloration often with several meters or more of soil developed in situ. In the more mafic lava flows of this formation the dominant ferromagnesian mineral, olivine, has been mostly to completely converted to iddingsite as a result of relatively low temperature oxidation and hydration. The most abundant mineral present in these lavas is the nonferromagnesian mineral plagioclase feldspar. It is often unaltered but in numerous locations it is cut by small cracks and veins filled by sericite mica and calcite plus a little quartz (?).

Where the ash flow tuff member of the Heppsie Formation has been exposed by erosion and due to its inherent structural weakness, numerous landslides and slumps

have their beginnings at or near the contact between younger more massive basaltic lava flows and the older poorly welded ash flow tuffs. In the adjacent Robinson Butte quadrangle (see Figure 1) several good examples of arcuate headwall slump structures can be easily found in the canyon of the South Fork of the Little Butte Creek. As one road-building civil engineer said to me (S.A.M.) two summers ago while working on a bridge replacement project across the South Fork, "Nothing in this valley is actually stable. Everything is moving downhill pretty fast, particularly in the spring as the winter snows melt." Also, numerous excellent examples of landslides / slump development can be found in the Klamath River canyon along the California – Oregon border beginning just below the John C. Boyle Powerhouse at what is known as the Caldera Rapid. In this region a very similar layering of geologic rock units occurs as in the South Fork of the Little Butte Creek: older siliceous poorly welded tuffaceous material is overlain by younger thick occurrences of mafic lava flows. It is likely that the same mechanism of slope failure is occurring in both areas, giving rise to landslides / slumps that set the stage for rapids to form in the streams that are flowing through the valleys. A half-dozen age dates for various Heppsie Formation samples, both whole rock K-Ar and $^{40}\text{Ar}/^{39}\text{Ar}$ ages, are available. However, none are from the minute exposures in the BMQ. The measured ages range from 21.73 ± 0.14 Ma to 19.93 ± 0.07 Ma old.

Acknowledgments I thank Isaac Weaver for his on-going help and support in bringing this geologic map to closure. His computer skills, particularly with regard to MapInfo and Adobe Illustrator, have been particularly valuable. I also thank Karen Mertzman for all her efforts in the x-ray lab, carefully preparing and analyzing countless samples on my behalf. Generous grants from the W. M. Keck Foundation supported fieldwork in the Brown Mountain quadrangle in 1991 and 1994. Support from the NSF and Franklin and Marshall College to facilitate the operation of the XRF laboratory in the Earth and Environment Department is greatly appreciated.

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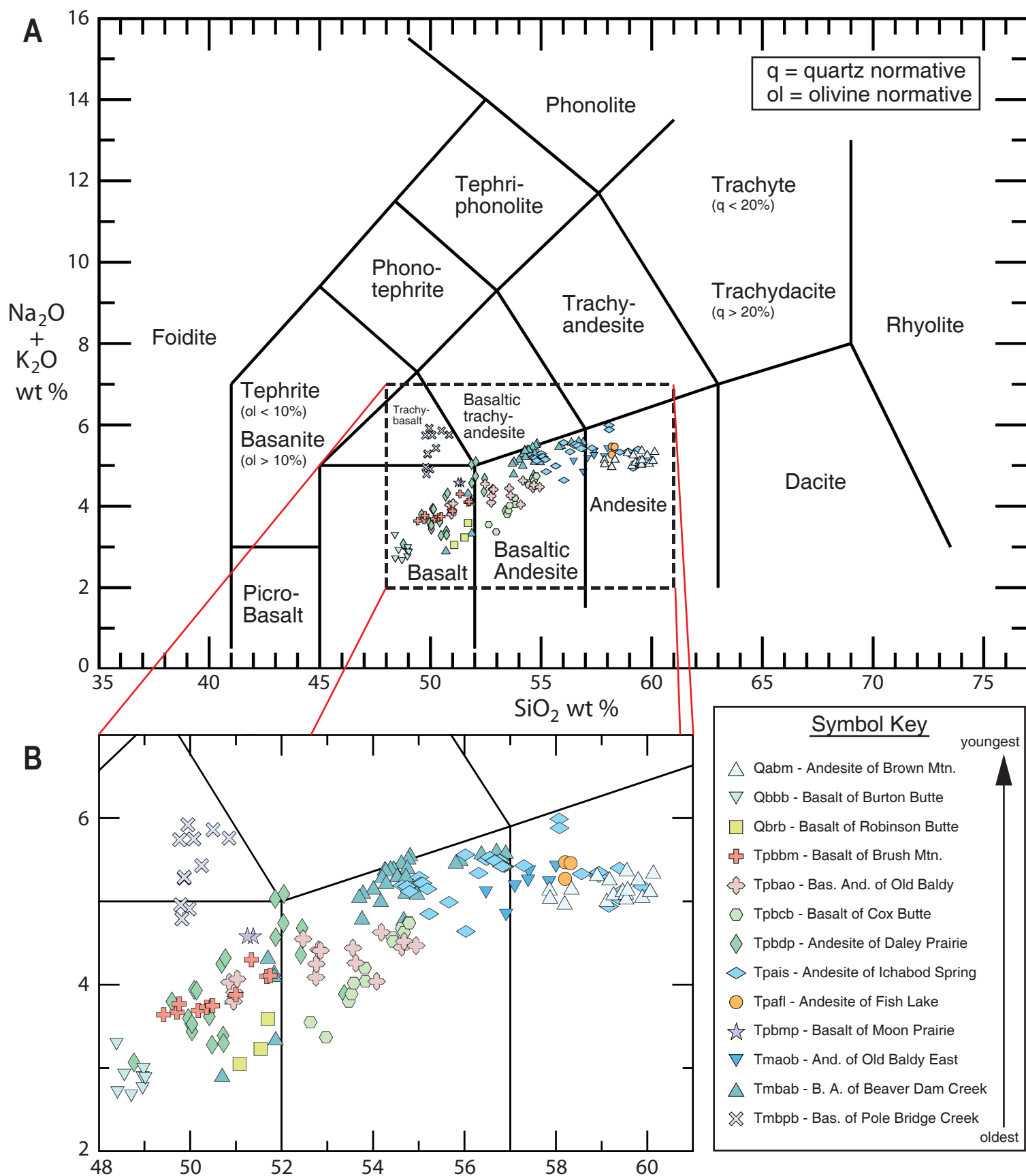


Figure 2. IUGS (International Union of Geological Sciences) classification system for volcanic rocks, which is based on total alkali (Na₂O + K₂O) vs. silica (SiO₂) content, with the data from analyzed Brown Mountain quadrangle samples (Table 1) superimposed (see Le Maitre, 2002).

Table 1 (page 1). Whole rock chemical data and Potassium-Argon (K-Ar) ages ^a indicates an Argon-Argon age) for the samples from the Brown Mountain quadrangle, Jackson and Klamath Counties, Oregon. The major element oxides are presented in weight percent and the trace elements are reported in parts per million (ppm). The chemical data are X-ray fluorescence (XRF) analyses and were measured in the X-ray laboratory of the Department of Earth and Environment, Franklin and Marshall College, Lancaster, Pennsylvania. The UTM coordinate values are according to the UTM Zone 10 (NAD 27 for US) projection. All UTM coordinates have been rounded to the nearest 10 m. The 1/4 of 1/4, 1/4, Section (Sec.), Township (T.), and Range (R.) columns are location descriptors of the Public Land Survey System (PLSS). In the Lithology column (Lith.), TrB = Trachybasalt (SiO₂ = 45-52%; Na₂O+K₂O ≥ 5.0%), B = Basalt (SiO₂ = 45-52%), BA = Basaltic Andesite (SiO₂ = 52-57%), and A = Andesite (SiO₂ = 57-63%). See Figure 2 and Le Maitre (2002) for details regarding lithological classification. Please consult the detailed write-up or the geologic map for the full unit names. Please note that this table is 3 pages long.

Map	Sample	K-Ar Age ^a	1/4	1/4 Sec.	T.	R.	UTME	UTMN	Unit	Lith.	SiO ₂	TiO ₂	Al ₂ O ₃	Fe ₂ O ₃	FeO	MnO	MgO	CaO	Na ₂ O	K ₂ O	P ₂ O ₅	LOI	Total	Fe ₂ O ₃ T	Rb	Sr	Y	Zr	V	Ni	Cr	Nb	Ga	Cu	Zn	Co	Ba	La	Ce	U	Th	Sc	Pb	Yb	Be
1	RS94-131	6.14 ± 0.15	SW	SE	20	37	4	551610	468680	TrB	48.77	1.07	15.56	5.41	3.10	0.14	7.49	9.70	3.35	2.39	0.64	0.74	99.36	8.86	22.4	2246	19.2	140	210	115	231	2.8	18.4	101	78	33	1925	41.3	103.7	4.5	8.2	36	7.4	1.8	1.5
2	99-5	4.35 ± 0.04 ^a	NE	SE	29	37	4	552300	468530	TrB	49.08	1.17	15.83	2.68	5.82	0.15	7.87	9.59	3.60	1.31	0.58	0.87	99.45	9.23	6.3	973	17.2	149	233	77	236	10.6	22.6	78	31	725	34	63	2.6	9.7	20	4	—	—	
3	99-17	—	SW	SE	29	37	4	551730	468520	TrB	49.87	1.09	15.09	3.68	5.55	0.14	9.14	9.39	3.80	1.16	0.75	0.86	99.44	9.23	8.3	1177	14.2	136	231	142	369	16.8	22.4	85	38	349	35	63	2.6	9.7	20	4	—	—	
4	00-51	—	SW	SE	29	37	4	551770	468550	TrB	49.82	1.13	15.20	5.76	3.57	0.14	8.77	8.51	3.61	1.18	0.74	1.08	99.51	9.23	12.4	1208	13	141	203	171	379	13.0	21.3	77	109	36	445	34	61	2.2	4.9	21	4	—	—
5	00-38	—	SW	SE	20	37	4	551740	468620	TrB	49.87	1.06	15.21	5.84	2.98	0.14	7.46	9.55	2.99	2.29	0.63	2.14	99.56	8.82	28.0	2163	24.2	137	176	101	207	8.8	17.5	85	86	33	2007	50	114	9.6	16	13	—	—	
6	00-39	—	NW	NE	29	37	4	551710	468640	TrB	49.87	1.07	15.37	7.27	1.62	0.15	6.90	9.03	2.98	2.31	0.62	2.41	99.60	9.07	29.9	1984	23.9	139	157	98	200	8.6	17.5	85	86	33	2030	45	106	2.3	9.9	20	11	—	—
7	00-35	—	NW	SE	20	37	4	551770	468710	TrB	49.85	1.07	15.37	7.27	1.62	0.15	6.90	9.03	2.98	2.31	0.62	2.41	99.60	9.07	29.9	1984	23.9	139	157	98	200	8.6	17.5	85	86	33	2030	45	106	2.3	9.9	20	11	—	—
8	02-46	—	NW	SE	20	37	4	551770	468730	TrB	50.08	1.15	15.16	5.69	2.72	0.13	7.62	7.43	3.47	2.32	0.71	3.27	99.71	8.71	32.3	1175	14.8	165	168	123	243	11.7	20.9	76	105	35	1421	49	109	2.6	11.3	14	10	—	—
9	00-40	—	NW	SE	29	37	4	551590	468630	TrB	50.25	1.07	15.58	7.23	1.68	0.14	7.03	9.25	3.36	2.07	0.57	1.32	99.55	9.10	18.6	2218	22.5	133	166	93	187	8.4	17.6	91	100	32	860	43	83	2.4	9.5	16	6	—	—
10	00-36	—	NW	SE	20	37	4	551700	468650	TrB	50.30	1.15	14.98	5.94	2.65	0.13	8.23	8.24	3.70	2.16	0.67	0.99	99.54	9.11	29.7	1502	15.8	178	176	164	263	10.7	19.9	105	98	36	983	43	91	1.9	9.3	17	10	—	—
11	00-37	—	SW	SE	20	37	4	551800	468620	TrB	50.85	1.12	14.67	5.80	2.91	0.13	8.53	8.09	3.65	2.11	0.65	0.85	99.56	9.03	29.1	1463	16	175	167	218	339	10.6	19.7	95	97	39	1045	38	86	2.3	10.0	15	9	—	—
12	91-44	4.60 ± 0.34	SW	NW	4	38	4	552800	468680	TrB	50.70	0.68	16.10	2.40	5.90	0.15	9.11	10.68	2.40	0.15	0.81	0.99	99.67	8.96	5.3	566	16.3	32	247	127	482	2.6	18.2	70	64	39	181	6	13	1.5	4.0	3.2	6.6	1.2	1.3
13	91-43	5.82 ± 0.17	SE	NW	4	38	4	552820	4682570	TrB	54.01	1.14	17.10	5.60	3.68	0.12	3.50	7.26	3.85	1.29	0.50	1.57	99.62	9.69	12.1	811	37.4	130	192	25	63	7.7	21.8	34	79	26	513	23	49	1.6	3.2	19	5.7	2.3	1.7
14	91-46	5.08 ± 0.13	SE	NW	4	38	4	552140	4684400	TrB	54.00	1.15	17.75	3.52	4.50	0.14	4.38	7.87	4.10	1.18	0.50	0.76	100.05	8.92	11.1	920	33.5	118	194	27	64	8.0	22.7	36	71	21	497	22	49	1.7	4.1	21	4.9	2.9	1.8
15	HP94-32	—	NE	SW	10	38	4	554640	4680600	TrB	50.92	1.06	17.14	0.78	7.25	0.16	7.26	9.33	3.17	0.65	0.25	1.03	99.00	8.94	8.1	588	18.9	83	171	73	307	4.7	17.8	75	62	29	308	10.9	29.6	1.4	1.9	22.9	6.8	—	—
16	HP94-22	—	NE	SW	10	38	4	554560	4680350	TrB	50.97	1.05	17.59	4.37	4.21	0.15	7.07	8.76	3.67	0.04	0.20	0.50	99.60	9.05	7.3	762	12.7	147	199	171	323	3.1	17.4	76	76	35	307	9.5	29.2	1.5	1.6	31.8	6.1	—	—
17	HP94-43	—	NE	NW	9	38	4	552580	4681410	TrB	51.70	0.82	17.57	4.37	4.21	0.15	7.07	8.76	3.67	0.04	0.20	0.50	99.60	9.05	7.3	762	12.7	147	199	171	323	3.1	17.4	76	76	35	307	9.5	29.2	1.5	1.6	31.8	6.1	—	—
18	00-54	—	SW	SE	5	38	4	551920	4687900	TrB	51.84	0.85	17.53	3.63	5.06	0.14	7.15	8.42	3.46	0.67	0.21	0.62	99.58	9.25	8.6	804	15.8	61	218	146	276	2.8	18.6	85	80	38	317	12	23	<0.5	<0.5	25	4	—	—
19	01-41	—	SE	SW	4	38	4	552830	4681900	TrB	51.84	0.86	16.64	4.14	0.74	0.14	7.59	8.75	3.41	0.68	0.19	0.58	99.52	9.36	8.7	825	16.1	31	218	128	288	2.5	18.2	55	71	37	336	8	18	1.1	—	24	—	—	—
20	99-22	—	NW	NW	33	37	4	552650	4684800	TrB	51.87	0.85	15.98	1.05	7.12	0.14	9.66	9.25	2.89	0.44	0.13	0.74	99.92	8.95	11.8	174	139	142	114	513	417	3.7	18.9	75	68	34	151	6	17	0.9	3.6	24	3	—	—
21	HP94-21	—	NW	NE	9	38	4	552650	4684800	TrB	51.87	0.85	15.98	1.05	7.12	0.14	9.66	9.25	2.89	0.44	0.13	0.74	99.92	8.95	11.8	174	139	142	114	513	417	3.7	18.9	75	68	34	151	6	17	0.9	3.6	24	3	—	—
22	HP94-7D	—	NW	SE	9	38	4	552030	4680350	TrB	51.70	0.82	17.57	4.37	4.21	0.15	7.07	8.76	3.67	0.04	0.20	0.50	99.60	9.05	7.3	762	12.7	147	199	171	323	3.1	17.4	76	76	35	307	9.5	29.2	1.5	1.6	31.8	6.1	—	—
23	02-47	—	NE	SE	20	37	4	552030	4680350	TrB	51.70	0.82	17.57	4.37	4.21	0.15	7.07	8.76	3.67	0.04	0.20	0.50	99.60	9.05	7.3	762	12.7	147	199	171	323	3.1	17.4	76	76	35	307	9.5	29.2	1.5	1.6	31.8	6.1	—	—
24	HP94-9	—	NE	SE	20	37	4	552030	4680350	TrB	51.70	0.82	17.57	4.37	4.21	0.15	7.07	8.76	3.67	0.04	0.20	0.50	99.60	9.05	7.3	762	12.7	147	199	171	323	3.1	17.4	76	76	35	307	9.5	29.2	1.5	1.6	31.8	6.1	—	—
25	HP94-23	—	NW	SE	10	38	4	552030	4680350	TrB	51.70	0.82	17.57	4.37	4.21	0.15	7.07	8.76	3.67	0.04	0.20	0.50	99.60	9.05	7.3	762	12.7	147	199	171	323	3.1	17.4	76	76	35	307	9.5	29.2	1.5	1.6	31.8	6.1	—	—
26	99-23	—	NW	SE	10	38	4	552030	4680350	TrB	51.70	0.82	17.57	4.37	4.21	0.15	7.07	8.76	3.67	0.04	0.20	0.50	99.60	9.05	7.3	762	12.7	147	199	171	323	3.1	17.4	76	76	35	307	9.5	29.2	1.5	1.6	31.8	6.1	—	—
27	99-4	—	NE	SE	2																																								

Table 1 (page 2).

Map no.	Sample no.	K-Ar Age Ma	1/4 Sec. of 1/4	T. (S.)	R. UTM E	UTM N	Unit	Lith.	SiO ₂	TiO ₂	Al ₂ O ₃	Fe ₂ O ₃	FeO	MnO	MgO	CaO	Na ₂ O	K ₂ O	P ₂ O ₅	LOI (%)	Total Fe ₂ O ₃ (%)	Rb	Sr	Y	Zr	V	Ni	Cr	Nb	Ga	Cu	Zn	Co	Ba	La	Ce	U	Th	Sc	Pb	Yb	Be			
57	91-68	3.38 ± 0.06	SE	NE 24	37	4	558550	4887570	Trails	BA	58.07	0.84	17.54	3.02	3.82	0.12	3.40	6.56	4.67	1.32	0.18	0.86	100.40	7.27	23.6	640	26.6	98	158	17	48	4.2	21.6	57	69	23	455	16	31	2.4	3.4	17	7.2	1.9	1.7
58	91-67	—	SE	NE 23	37	4	558910	4887000	Trails	BA	54.64	0.96	18.38	3.25	4.57	0.14	3.95	7.32	4.24	0.95	0.23	8.33	99.36	8.33	13.4	668	32.6	92	150	33	53	4.9	20.3	31	71	21	386	13	25	1.8	2.9	20	6	—	—
59	99-7	—	SW	NE 21	37	4	552320	4887500	Trails	BA	54.74	0.99	18.13	2.93	5.59	0.15	3.97	7.26	4.22	0.96	0.24	1.09	99.67	8.54	11.8	647	22.2	84	195	35	49	4.7	22.2	53	74	24	417	12	29	1.1	4.6	19.4	6	—	—
60	99-19	—	SW	NW 21	37	4	552320	4887640	Trails	BA	54.80	1.01	18.10	2.96	5.20	0.15	4.21	7.24	4.17	0.95	0.24	1.00	100.05	8.76	12.1	644	26.2	84	201	39	39	4.9	22.0	112	73	23	374	15	22	1.6	4.7	19	4	—	—
61	00-11	—	NE	NE 22	37	4	555430	4887730	Trails	BA	55.01	1.02	17.88	5.20	3.26	0.14	3.96	7.07	4.22	1.07	0.26	1.04	100.13	8.82	13.8	643	29.4	92	199	45	24	6.1	18.0	97	79	26	420	13	26	0.7	1.0	22	4	—	—
62	00-13	—	SW	NE 23	37	4	556480	4886600	Trails	BA	55.04	1.03	18.04	4.15	4.31	0.14	3.97	7.22	4.21	1.01	0.24	1.15	100.49	8.84	12.3	669	40	95	186	44	19	5.7	19.0	124	74	25	412	5.7	18.9	<0.5	1.0	22	—	—	
63	99-20	—	SW	NW 22	37	4	554190	4887500	Trails	BA	55.08	0.98	18.52	2.95	4.71	0.14	3.82	7.20	4.21	0.97	0.23	1.14	100.01	8.18	12.7	650	27.4	85	186	40	48	4.5	22.2	63	70	22	393	18	25	1.1	4.7	19	5	—	—
64	91-72	—	NE	NE 23	37	4	558820	4887500	Trails	BA	55.19	1.02	17.68	3.72	4.47	0.14	4.20	7.27	4.14	1.01	0.23	0.97	99.64	8.69	17.4	641	27.4	80	194	31	74	3.2	19.6	55	76	24	401	12	26	0.7	0.6	20	—	—	
65	91-64	—	NE	NW 28	37	4	553110	4886270	Trails	BA	55.23	1.00	17.84	3.43	4.62	0.13	4.17	7.27	3.78	1.07	0.21	0.80	100.00	8.56	19.1	628	33.4	80	198	21	64	4.0	18.1	67	74	24	409	14	28	—	1.7	22	—	—	
66	84-17	—	SW	SW 14	37	4	555800	4886480	Trails	BA	55.66	1.00	17.76	3.86	4.21	0.15	4.11	7.03	3.97	1.02	0.26	0.81	99.64	8.58	16.3	632	20	91	186	35	28	4.3	18.7	47	74	24	420	9	31	2.4	1.2	23	4	—	—
67	91-70	—	SE	NE 16	37	4	553880	4886220	Trails	BA	56.02	0.93	17.43	4.45	3.00	0.13	4.09	7.05	4.12	1.44	0.25	0.61	99.52	7.78	22.8	595	25.5	140	162	46	121	6.8	19.1	82	70	21	443	11	21	1.2	1.4	23	6	—	—
68	00-24	—	NE	NW 32	37	5	560970	4884850	Trails	BA	56.03	0.75	17.04	2.69	4.06	0.12	6.32	7.42	3.61	1.03	0.24	0.76	100.07	7.20	9.8	903	20.9	118	131	160	281	60	18.6	63	66	28	474	17	36	1.4	1.2	18	4	—	—
69	91-74	—	SW	NW 16	37	5	559090	4886790	Trails	BA	56.31	0.79	18.74	2.28	4.51	0.13	3.53	6.87	4.16	1.16	0.16	0.85	99.53	7.29	21.2	654	25.2	140	149	35	76	5.1	20.1	119	129	16	421	9	29	1.5	3.4	19	9	—	—
70	99-3	—	SW	NW 16	37	4	552590	4886300	Trails	BA	56.48	1.00	17.10	3.12	4.42	0.13	4.13	6.85	4.00	1.50	0.26	1.14	100.35	8.03	22.9	551	24.2	140	161	41	75	6.3	21.9	112	73	500	15	23	1.4	5.1	18	10	—	—	
71	00-34	—	SE	SE 16	37	4	553690	4886500	Trails	BA	56.55	1.01	17.06	2.85	4.76	0.13	4.14	6.89	4.00	1.53	0.26	0.87	100.05	8.14	25.0	605	32.4	145	169	47	71	6.6	18.8	83	74	24	563	18	37	<0.5	1.6	22	—	—	
72	91-62	—	NW	NE 21	37	4	553220	4886090	Trails	BA	56.67	1.03	16.67	3.70	3.77	0.12	4.40	3.98	1.51	0.20	0.59	99.68	7.89	25.4	586	29.2	137	185	37	83	4.3	18.8	95	78	25	482	15	37	<0.5	1.6	22	—	—		
73	H031-53	—	NE	NW 10	37	4	554620	4886940	Trails	BA	56.83	1.03	17.06	4.15	3.50	0.13	3.81	6.81	3.85	1.57	0.26	1.04	100.04	8.04	21	485	39	138	182	29	72	—	—	—	—	—	—	—	—	—	—	—	—		
74	91-71	—	SE	NE 16	37	4	553880	4886220	Trails	BA	56.87	1.03	17.18	4.60	3.36	0.12	3.88	6.70	3.88	1.54	0.24	1.18	100.58	8.33	25	602	27	145	182	50	80	—	—	—	—	—	—	—	—	—	—	—	—	—	
75	00-30	—	NE	SW 36	37	4	557950	4883980	Trails	A	57.30	0.87	17.48	3.51	3.05	0.11	3.75	6.80	4.01	1.42	0.25	1.09	99.64	6.90	19.5	1216	24.7	130	144	53	63	6.4	19.2	59	68	23	573	20	45	1.0	2.5	16	5	—	—
76	00-10	—	SE	NE 15	37	4	555200	4886050	Trails	A	56.08	0.85	17.19	3.53	3.09	0.11	3.49	6.89	4.27	1.61	0.20	1.75	100.26	6.86	26.9	676	23.6	127	144	45	46	6.3	20.0	68	77	21	536	17	31	1.2	3.1	16	10	—	—
77	91-81	—	SE	SE 35	37	4	557050	4883690	Trails	A	58.56	0.89	16.97	3.58	2.98	0.11	3.49	6.61	3.86	1.47	0.24	1.23	99.99	6.89	21.4	1152	29.1	111	147	36	74	4.2	19.8	91	70	20	504	30	50	1.6	2.1	16	—	—	
78	00-14	—	NE	NW 25	37	4	557620	4886180	Trails	A	58.94	0.72	17.90	3.69	2.53	0.10	3.14	6.04	4.10	1.20	0.20	1.70	100.26	6.50	11.5	809	16.3	101	127	32	31	5.6	21.3	31	59	18	388	13	26	<0.5	1.1	16	6	—	—
79	00-8	—	NE	NE 10	37	4	5560150	4886490	Trails	A	58.94	0.71	17.94	5.11	1.25	0.10	3.36	6.45	4.16	1.19	0.19	0.88	100.28	6.50	11.5	816	21.2	100	136	34	29	6.0	21.6	30	61	19	430	15	30	0.6	2.4	15	6	—	—
80	84-18	—	NE	NE 25	37	4	556330	4886490	Trails	A	59.16	0.69	17.95	2.94	2.68	0.10	3.49	6.14	3.82	1.12	0.20	1.24	99.53	5.92	13.0	840	14.1	100	130	30	33	4.1	19.2	22	55	14	403	11	25	0.7	0.6	14	2	—	—
81	00-9	—	NE	NW 27	37	4	558400	4886330	Trails	A	59.39	0.71	17.94	3.32	2.62	0.10	3.34	6.51	4.18	1.22	0.20	0.81	100.34	5.23	12.8	855	16.7	101	130	36	32	5.4	21.5	32	62	17	411	10	21	0.6	1.5	16	5	—	—
82	91-88	—	SW	SW 32	37	5	560890	4883450	Trails	A	59.39	0.72	17.97	2.76	3.14	0.10	3.34	6.54	3.92	1.11	0.18	0.97	100.14	6.25	12.4	820	18.7	82	160	23	49	3.3	20.0	40	62	17	388	13	26	0.5	0.8	16	—	—	
83	91-66	—	NE	SE 26	37	4	556880	4886870	Trails	A	59.42	0.72	17.87	3.30	2.77	0.10	3.29	6.11	3.89	1.23	0.18	1.21	100.09	6.38	14.5	797	26.4	88	143	24	50	3.2	19.6	37	63	18	452	16	22	0.5	1.1	12	3	—	—
84	00-7	—	SE	SW 29	37	5	561040	4886230	Trails	A	59.48	0.59	18.28	2.36	3.15	0.09	2.88	6.11	3.89	1.23	0.18	1.21	100.11	5.86	8.8	780	16.9	81	110	18	16	4.5	22.0	53	53	15	353	12	22	0.5	1.1	12	3	—	—
85	91-39	2.92 ± 0.07	NE	NE 16	37	5	560490	4883280	Tbpd	B	50.01	0.95	16.67	2.64	6.07	0.15	9.88	2.95	0.58	0.31	0.39	99.42	9.39	5.1	737	23.8	80	198	151	356	3.4	18.5	48	69	42	277	15	38	1.8	2.1	3.8	1.9	1.4		
86	91-76	2.77 ± 0.05	SE	SE 30	37	5	560490	4883300	Tbpd	B	50.01	0.95	16.67	2.64	6.07	0.15	9.88	2.95	0.58	0.31	0.39	99.42	9.39	5.1	737	23.8	80	198	151	356	3.4	18.5	48	69	42	277	15	38	1.8	2.1	3.8	1.9	1.4		
87	94-21A	2.96 ± 0.11	NW	NE 11	38	4	554840	4887640	Tbpd	B	50.04	0.97	16.87	3.63	5.01	0.15	10.63	9.82	2.67	0.77	0.33	1.07	99.69	9.39	6.5	709	24.2	97	202	101	284	69	173	54	69	35	483	18	39	<0.5	0.6	31	6	—	—
88	94-20	3.16 ± 0.12	NW	NE 10	38	4	554840	4887640	Tbpd	B	50.04	1.00	17.32	5.07	3.89	0.15	8.01	9.82	2.93	0.80	0.29	1.34	100.46	9.39	6.5	709	24.2	97	202	101	284	69	173	54	69	35									

Table 1 (page 3).

no.	Sample no.	K-Ar Age Ma	1/4 of 1/4	1/4 Sec. T. (S.)	R. UTM E m	UTM N m	Unit	Lith.	SiO ₂	TiO ₂	Al ₂ O ₃	Fe ₂ O ₃	FeO	MnO	MgO	CaO	Na ₂ O	K ₂ O	P ₂ O ₅	LOI (%)	Total (%)	Fe ₂ O ₃ Tb	Sr	Y	Zr	V	Ni	Cr	Nb	Ga	Cu	Zn	Co	Ba	La	U	Th	Sc	Pb	Yb	Be					
121	AL94-57	--	SE	NW	12	38	4	557200	4881220	Tpbaio	B	50.86	1.05	17.72	3.55	5.85	0.15	6.89	8.94	3.48	0.55	0.21	1.07	100.32	9.52	7.9	563	21.4	78	254	133	213	4.7	18.0	33	75	34	239	9.3	22.3	0.8	1.1	25.8	6.9	1.5	1.2
122	AL94-58	--	SE	NW	11	38	4	557200	4881590	Tpbaio	B	50.95	1.03	17.60	2.12	6.66	0.16	7.72	8.63	3.25	0.55	0.23	1.14	100.04	9.52	7.9	511	19	83	182	141	309	4.8	17.5	67	71	33	314	8.4	25.3	--	2.3	26.5	7.5	--	--
123	AL94-43	--	SE	NE	13	38	4	556860	4870300	Tpbaio	B	50.97	1.05	17.59	1.30	7.76	0.15	6.97	9.45	3.35	0.56	0.25	0.36	100.23	9.94	7.3	546	18.9	75	228	109	204	4.9	17.8	11	80	33	248	7.8	20.1	1.0	1.3	26.1	5.1	1.5	1.2
124	AL94-48	--	SE	NE	12	38	4	556860	4880120	Tpbaio	B	51.03	1.07	17.64	3.42	5.60	0.15	7.12	8.99	3.52	0.55	0.23	0.36	99.68	9.62	6.3	575	19	85	190	119	191	4.7	18.1	43	72	29	240	6.9	24.6	0.3	1.8	24.8	4.8	1.0	1.1
125	B594-3	--	SW	NW	19	38	5	559110	4877640	Tpbaio	BA	52.47	1.13	17.47	1.73	7.10	0.15	6.34	7.99	3.78	0.77	0.35	0.16	99.44	9.62	12	581	22.9	110	175	101	230	6.9	17.7	68	82	31	354	12.7	33	0.6	0.9	21.7	6.8	0.9	2
126	AL94-53	--	SW	SW	13	38	4	557640	4876530	Tpbaio	BA	52.81	0.93	17.64	1.65	6.49	0.14	5.44	8.25	3.53	0.72	0.34	1.05	99.29	8.86	12.1	542	18.5	96	174	80	155	6.4	19.4	60	69	25	356	11.6	29.8	0.4	1.9	22.7	7.7	1.6	1.2
127	AL94-53	--	SW	SW	13	38	4	557640	4876530	Tpbaio	BA	52.81	0.93	17.64	1.65	6.49	0.14	5.44	8.25	3.53	0.72	0.34	1.05	99.29	8.86	12.1	542	18.5	96	174	80	155	6.4	19.4	60	69	25	356	11.6	29.8	0.4	1.9	22.7	7.7	1.6	1.2
128	HP94-37	--	NE	SW	11	38	4	556180	4880590	Tpbaio	BA	52.85	1.03	17.83	2.62	5.69	0.15	5.66	8.21	3.68	0.73	0.34	0.85	99.30	8.81	9	634	20.2	96	174	80	155	6.4	19.4	60	69	25	356	11.6	29.8	0.4	1.9	22.7	7.7	1.6	1.2
129	AL94-36A	--	NE	SW	11	38	4	556180	4880590	Tpbaio	BA	52.85	1.03	17.83	2.62	5.69	0.15	5.66	8.21	3.68	0.73	0.34	0.85	99.30	8.81	9	634	20.2	96	174	80	155	6.4	19.4	60	69	25	356	11.6	29.8	0.4	1.9	22.7	7.7	1.6	1.2
130	AL94-36A	--	NE	SW	11	38	4	556180	4880590	Tpbaio	BA	52.85	1.03	17.83	2.62	5.69	0.15	5.66	8.21	3.68	0.73	0.34	0.85	99.30	8.81	9	634	20.2	96	174	80	155	6.4	19.4	60	69	25	356	11.6	29.8	0.4	1.9	22.7	7.7	1.6	1.2
131	AL94-47	--	SW	NE	13	38	4	557540	4879830	Tpbaio	BA	54.18	0.89	18.14	1.51	6.55	0.14	5.42	8.01	3.57	0.69	0.24	0.56	100.40	8.79	12.7	589	20.7	77	173	61	133	5.7	19.7	71	69	23	418	17	28.3	0.7	1.3	22.7	8.2	1.9	1.2
132	B594-104	--	SW	NE	11	38	4	557000	4877800	Tpbaio	BA	54.18	0.89	18.14	1.51	6.55	0.14	5.42	8.01	3.57	0.69	0.24	0.56	100.40	8.79	12.7	589	20.7	77	173	61	133	5.7	19.7	71	69	23	418	17	28.3	0.7	1.3	22.7	8.2	1.9	1.2
133	HP94-57	--	SW	NE	11	38	4	556900	4877800	Tpbaio	BA	54.63	0.88	18.75	3.87	3.92	0.13	7.97	7.83	3.77	0.68	0.21	0.49	103.13	8.23	12	589	18.2	76	155	62	138	4.7	18.4	33	70	23	354	11.8	26.2	--	0.8	21	7.4	1.7	1
134	HP94-34	--	SW	NE	10	38	4	555900	4880220	Tpbaio	BA	54.68	0.86	17.94	1.64	5.69	0.14	5.23	8.01	3.77	0.75	0.26	0.66	99.56	7.96	11.9	649	19	89	146	146	146	146	146	146	146	146	146	146	146	146	146	146	146	146	146
135	AL94-34	--	SE	NW	23	38	4	556180	4877690	Tpbaio	B	49.71	1.20	18.09	1.39	7.79	0.16	6.74	10.25	3.20	0.47	0.20	1.26	100.46	10.05	6.8	469	30	68	229	57	166	3.1	20.4	66	69	31	246	12	27	2.5	3.8	32	3.8	2.7	1.2
136	AL94-34	--	SE	NW	23	38	4	556180	4877690	Tpbaio	B	49.71	1.20	18.09	1.39	7.79	0.16	6.74	10.25	3.20	0.47	0.20	1.26	100.46	10.05	6.8	469	30	68	229	57	166	3.1	20.4	66	69	31	246	12	27	2.5	3.8	32	3.8	2.7	1.2
137	HP94-62	--	SW	NE	23	38	4	556950	4878270	Tpbaio	B	49.42	1.29	17.77	2.21	7.75	0.17	6.81	10.50	3.21	0.43	0.18	0.87	100.61	10.82	6.8	458	24.2	65	217	62	163	4.6	19.4	73	64	30	199	8.1	21.1	0.7	--	29.8	6.3	1.7	1.2
138	HP94-13	--	SW	NE	15	38	4	556040	4879280	Tpbaio	B	49.76	1.23	17.74	2.00	7.35	0.17	6.56	10.38	3.30	0.47	0.21	0.91	100.08	10.17	6.1	487	26	78	183	57	133	4.6	19.2	86	61	27	232	7.7	16	1.5	2.0	28	6.6	--	--
139	HP94-42	--	NE	NW	15	38	4	554070	4878350	Tpbaio	B	50.18	1.08	17.45	1.54	7.36	0.17	7.52	10.14	3.04	0.65	0.24	1.11	100.48	9.72	6.3	640	17.3	69	150	68	308	4.3	16.8	73	60	23	281	8.5	25.6	1.4	1.8	20.9	6.7	--	--
140	HP94-46	--	NE	NW	15	38	4	554810	4879980	Tpbaio	B	50.36	1.05	17.63	1.57	7.73	0.16	6.94	9.78	3.21	0.52	0.24	0.91	100.10	10.16	4.9	609	19	83	194	76	324	4.9	18.0	79	66	28	315	12.9	33.4	0.8	2.0	26.2	7.7	--	--
141	99-24	--	NE	NW	20	38	4	552400	4876340	Tpbaio	B	50.36	1.06	17.49	1.27	7.35	0.16	7.14	9.91	3.17	0.56	0.23	0.86	99.96	9.44	5.7	605	19.2	76	177	317	4.6	18.7	75	63	28	308	12.9	27.4	1.0	3.1	27.4	7.5	--	--	
142	HP94-49	--	NE	NW	20	38	4	552340	4876280	Tpbaio	B	50.48	1.07	17.38	1.00	7.94	0.16	7.23	9.61	3.16	0.59	0.23	1.09	99.94	9.42	4.8	597	22.6	82	219	310	5.3	20.7	82	67	33	289	12	23	1.7	4.6	25	5	--	--	
143	94-17	--	SE	NW	23	38	4	553680	4879120	Tpbaio	B	50.49	1.04	17.54	1.12	7.35	0.16	7.01	9.70	3.17	0.58	0.24	0.83	99.23	9.29	5.2	614	21.3	89	192	316	4.7	18.5	74	66	28	283	16.1	23	2.2	2.2	25.2	5.7	--	--	
144	AL94-31	--	SW	NW	23	38	4	556950	4877670	Tpbaio	B	51.00	1.05	17.50	3.02	6.13	0.15	7.87	8.37	3.35	0.53	0.26	0.98	100.21	9.83	9.4	562	25	91	192	337	366	3.3	17.6	33	76	30	320	--	30	<5	2.2	28	--	2.8	1.2
145	AL94-36E	--	NE	SW	14	38	4	556770	4878240	Tpbaio	B	51.34	0.90	17.14	2.46	6.38	0.15	7.97	8.02	3.64	0.66	0.31	0.74	99.71	9.55	9.9	638	20.7	107	127	364	5.9	17.0	42	76	31	286	16.6	36.9	0.8	2.5	23.3	5.1	1.5	1.2	
146	AL94-38	--	NE	SW	14	38	4	556250	4879050	Tpbaio	B	51.68	1.00	17.44	2.64	5.87	0.14	7.25	9.02	3.41	0.69	0.26	0.65	100.05	9.16	7.9	654	12	60	174	91	192	4.5	20.2	39	71	30	303	7.2	36.3	0.5	2.1	24.7	6.7	2.4	1.2
147	AL94-16	0.40 +/- 0.30	SE	SE	8	37	4	552200	4869880	Qbba	B	51.75	0.95	18.13	9.75	--	0.15	6.28	7.60	3.44	0.67	0.33	--	99.06	9.75	10	628	26	122	163	150	319	6.9	17.7	25	80	31	328	19.4	38.4	0.7	2.7	25	5.3	1.7	1.2
148	99-2	--	SE	SW	16	37	4	552930	4868440	Qbba	B	51.70	0.95	18.05	2.38	5.88	0.14	6.84	9.98	3.19	0.40	0.13	0.88	100.47	8.86	4.9	811	21.1	53	210	46	146	2.6	20.4	67	58	33	167	12	23	0.5	3.0	26	3.2	1.7	1
149	99-1	--	NE	NW	21	37	4	553000	4868010	Qbba	B	51.08	0.85	16.41	1.28	6.51	0.14	9.37	9.81	2.66	0.39	0.16	1.51	100.17	8.81	1.8	897	16.6	75	210	172	427	4.3	19.0	75	68	37	193	14	31	1.2	4.3	28	3	--	--
150	00-58	--	SE	SW	34	37	4	554680	4883320	Qbba	B	51.54	0.86	16.53	2.09	5.74	0.14	9.26	9.73	2.79	0.44	0.16	1.07	100.35	8.47	1.9	913	16.1	73	191	172	435	4.2	18.8	70	62	35	173	9	34	1.0	5.0	26	4	--	--
151	91-42	--	NE	SW	4	38	4	552700	4882380	Qbba	B	48.41	1.03	16.97	3.25	6.96	0.17	7.75	10.04	2.99	0.32	0.21	1.03	99.83	10.98	2.9	463	22.5	84	223	149	223														