OPEN FILE REPORT #0-80-2 RECONNAISSANCE GEOLOGIC MAP OF THE BELKNAP - FOLEY AREA, OREGON

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EXPLANATION

(Boundaries are approximate; statements are general; lithologies are based on hand specimen identification)

Qgf

Qmt

Qbh

<u>Glaciofluvial deposits</u>: Unconsolidated, locally stratified sand, gravel, silt, cobbles, and boulders up to 5 m in diameter, with lenses of varved silts and clays where glacial deposits predominate. Forms flat-lying to hummocky deposits on floors of stream valleys. Includes undifferentiated alluvial, terminal moraine, and glacial outwash deposits. Approximate equivalent to Qal (alluvium) of Wells and Peck (1961), Qal (alluvium) of Peck and others (1964), and Qgt (glacial till) and Qal (alluvium) of Jan (1967) (1967)

Moraine terrace deposits: Unconsolidated, unstratified sand, gravel, silt, cobbles, and boulders, with lenses of varved silts and clays. Forms blan-ket on lower portions of slopes of McKenzie River and Horse Creek valleys upstream from Rainbow. Composed of undifferentiated lateral moraine, glacial outwash, and talue deposits. glacial outwash, and talus deposits

Landslide deposits: Unconsolidated blocks of bed rock and debris moved Qis downslope via debris flows and slumping from undercutting by stream action. Material size ranges from silt- and clay-size particles to blocks 1/4 km in longest dimension. Prominent scarps are normally present at heads of slides. Partially equivalent to Qls (landslide) of Swanson and James (1975a)

Pleistocene to Holocene basalts of the High Cascades: Holocene, black, glassy, olivine-bearing, vesicular, basaltic flow tongues; flow structures common and pronounced; overlies recent glacial and fluvial deposits in stream valleys. Found on north fork of McKenzie River, Anderson Creek, and Lost Creek White Branch. Approximately equivalent to "Anderson Creek flows" and "Sims Butte flows" of Taylor (1968); partially equivalent to Qv (volcanic rocks of the High Cascades) of Callaghan and Buddington (1938), Qb (basalt) and QTba (basalts and andesites) of Wells and Peck (1961), QTv (volcanic rocks of the High Cascades . . . undivided) of Peck and others (1964), and T/Qgb (glomeroporphyritic basalt) of Jan (1967)

and James (1975a)

Dacite of Castle Rock: Upper Miocene, medium- to light-gray, dense, quartz-bearing dacitic rocks. Occurs at Castle Rock intruding flows of unit *Tmv*; Tindc includes related plug, flow rock, and sparse interflow pyroclastic rocks. Partially equivalent to Ta (volcanic rocks of the Western Cascades) of Callaghan and Buddington (1938), Tmua (andesite and basalt) of Wells and Peck (1961), and Tsa (Sardine Formation andesite) of Peck and others (1964)

Miocene voicanic rocks: Miocene basaltic, andesitic, and dacitic lava flows and autobreccias and, to a lesser extent, tuffs, lapilli tuffs, ash flows, Tmvd Tmv mudflows, debris flows, and minor epiclastic rocks. Dark-gray to black, dense, porphyritic, two-pyroxene andesites with large plagioclase phenocrysts predominate, but aphyric basaltic rocks are also common. Mafic minerals are generally altered to green clay minerals, with degree of alteration increas-ing toward base of unit. Flow contacts are readily identifiable in outcop, Tinvi and several unconformities are present within unit. Caps ridges west of Lookout Mountain and Cougar Reservoir and is dominant rock type east of Lookout Mountain and Cougar Reservoir and west of Horse Creek and north fork of McKenzie River. Limited outcrops underlie unit T_{Pv} near Scott Creek and eastern edge of Foley Ridge. Unit is defined on basis of lithology, stratigraphic position, and K/Ar dates and includes all similar rocks of Miocene age (unit Imde excluded). Several large intrusive and remnant volcanic plugs are mapped separately as unit *Imvi*, and basaltic and andesitic dikes in Blue River valley are mapped separately as unit *Imot*, and basartic and andestric dives in Blue River valley are mapped separately as unit *Imod*. Partially equiva-lent to Ta (volcanic rocks of the Western Cascades) of Callaghan and Buddington (1938) and "basalt" of Tuck (1927). Approximate time-strati-graphic equivalents of Tmua (andesite and basalt) of Wells and Peck (1961), Tsa (Sardine Formation andesite) of Peck and others (1964), Tsf (basaltic and basaltic andesite flows and breccias) of Jan (1967), "Sardine Formation" of Swanson and James (1975a), and "Sardine Formation" of Storch (1978)

Miocene dioritic rocks: Middle (?) Miocene, light-gray, aphanitic, equi-Tind granular, quartz-bearing diorite; crops out locally in Tidbits Creek drainage in Blue River mining district, intruding flows of lower Imo unit. Approxi-mately equivalent to Ti (intrusive igneous rocks) of Callaghan and Buddington (1938) and Tg (granodiorite) of Peck and others (1964)

and, to a lesser extent, black, dense, basaltic and andesitic autobreccias and flows. Pyroclastic rocks are generally highly sheared and massive in outcrop and altered to contain green, greenish-brown, and light-brown, low-grade metamorphic minerals. Lava flows are generally highly altered, with mafic minerals wholly altered to green and white clay minerals. Margins of flows are obscured in outcrop by numerous shear zones filled with clay, calcite, and green to white quartz and chalcedony. Unit is easily discernible from unit $\mathcal{I}mv$ in outcrop on basis of lithology, relative degree of alteration, and shearing. Unit is dominant rock type west of Lookout Mountain and Cougar Reservoir. Unit is defined on basis of lithology and stratigraphic position and includes all similar rocks lying below unit *Imv*. Partially equivalent to "andesite-agglomerate" of Tuck (1927) and Ta (volcanic rocks



GEOLOGIC SYMBOLS





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STATE OF OREGON DEPARTMENT OF GEOLOGY AND MINERAL INDUSTRIES 1005 State Office Building Portland, Oregon 97201

OPEN-FILE REPORT 0-80-2

PRELIMINARY GEOLOGY AND GEOTHERMAL RESOURCE POTENTIAL OF THE BELKNAP-FOLEY AREA, OREGON

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DISCLAIMER

This report has not been edited for complete conformity with Oregon Department of Geology and Mineral Industries standards. Data in this document are preliminary and are subject to change upon further verification.

CONTENTS

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INTRODUCTION
GEOLOGY
Introduction
Volcanic stratigraphy
Structural geology
Relations of structures to geothermal systems 9
GEOPHYSICS
WATERCHEMISTRY
GEOTHERMAL-GRADIENT AND HEAT-FLOW DATA
CONCLUSIONS AND RECOMMENDATIONS
BIBLIOGRAPHY OF THE BELKNAP-FOLEY AREA
APPENDIX A: Formulas used in calculations
APPENDIX B: Geothermal-gradient data

.

ILLUSTRATIONS

-

FIGURES

1. Map showing location of study area, Belknap-Foley area, Oregon \cdot . 2
2. Photo-lineament map of Belknap-Foley area 8
3. Total field aeromagnetic anomaly map of Belknap-Foley area 11
4. Complete Bouguer anomaly map of Belknap-Foley area 12
5. Residual gravity anomaly map of Belknap-Foley area 13
TABLES
1. Radiometric (K/Ar) ages of selected rocks, Belknap-Foley area 4
2. Bulk chemical composition of selected rocks, Belknap-Foley area 5
3. Spring and well chemistry, Belknap Foley area 16
4. Geothermetric calculations of minimum reservoir temperatures
for thermal waters, Belknap-Foley area
5. Geothermal-gradient data, Belknap-Foley area 22
MAPS (folded, in envelope)

plate I. Reconnaissance geology of the Belknap-Foley area, Oregon

INTRODUCTION

The Belknap-Foley area is located in the central Western Cascade Range of Oregon, approximately 80 km (50 mi) east of Eugene (Figure 1). Limits of the study area were arbitrarily assigned by U.S. Geological Survey (USGS) topographic map limits and natural breaks in the geology and topography (Plate I). This study, performed under U.S. Department of Energy (USDOE) Contract No. DE FC07-79ET27220, was undertaken to estimate the geothermal potential of the area by using various methods including compilation of existing data, reconnaissance geologic mapping, lineament analysis, well and spring geochemistry, and accrual of geothermal-gradient data.

Geographically, the study area is located in the rugged mountains surrounding the valley of the McKenzie River, which bisects the area in an eastwest direction. Total relief is approximately 1,000 m (3,300 ft) in the mountainous areas and approximately 30 m (100 ft) in the river valley.

GEOLOGY

Introduction

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The Belknap-Foley area is located at the eastern boundary of the Western Cascades geologic province in the Western Cascades-High Cascades transition zone. Quaternary and late Tertiary lavas and minor tuffs of the High Cascades province are in steep depositional contact with older Western Cascades rocks along this boundary, which appears to be the western margin of a major northsouth-trending High Cascades graben (Allen, 1966; Taylor, 1978, 1980). Because a number of thermal springs and preexisting gradient holes with high values are located along the margin of this graben, much of the mapping effort of this study was directed at carefully defining the nature of the High Cascades-Western Cascades geologic boundary.

- 1 -



Figure 1: Map showing location of study area.

2

The geology and all K/Ar radiometric ages (Table 1) are presented on the accompanying reconnaissance geologic map (Plate I), which was produced during the summer and fall of 1979 and 1980. Areal extent of geologic units was based on mapping and hand-specimen identification of rocks. Data were plotted on USGS topographic maps without the aid of aerial photographs.

Volcanic stratigraphy

From middle Tertiary to Quaternary time, volcanism in the area changed from silicic pyroclastic activity to eruption of increasingly mafic magmas (Table 2). This change in composition was reflected in higher percentages of lavas relative to tuffs. The oldest mappable unit (unit Tov on the geologic map) is composed of epiclastic volcanic sedimentary rocks, lithic-fragmentrich laharic dacite tuffs, and minor mafic lava flows. These rocks are probably Oligocene to early Miocene in age (Peck and others, 1964). In the Blue River mining district, the Oligocene rocks are locally intruded by Miocene quartzbearing dioritic stocks (unit *Imd*). The Miocene sequence (unit *Imv*) is dominated by highly phyric lavas, autobreccias, and mudflows with two-pyroxene andesite clasts, although lesser volumes of ash-flow, air-fall, and epiclastic tuffs as well as some basaltic flows occur locally. Several Miocene volcanic plugs and plug domes (unit Imvi) occur in the western part of the area, and basaltic to andesitic feeder dikes (unit *Imvd*) occur in the Blue River valley. The oldest dated rock assigned to the Miocene volcanic sequence in the map area is 19.91+1.94 m.y. old (McBirney and others, 1974). The youngest dated Imu sample is 6.2+0.2 m.y. old (Laursen and Hammond, 1978). The Miocene rocks are overlain by diktytaxitic to compact basaltic to basaltic-andesitic lavas and one small ash flow which cap most of the high ridges in the western part of the area. These Pliocene volcanic rocks (unit $T_{\mathcal{P}}v$) have a maximum K/Ar age of 8.39+0.36 m.y. (unpublished University of Utah Research Institute (UURI) K/Ar data, Evans and Foley, analysts) and a minimum age of 3.88+0.06 m.y.

- 3 -

Sample no.*	Location	Rock type	<u>Age</u> **	Stratigraphic unit
MS-254	122 ⁰ 07'30" 44 ⁰ 14'10"	Basalt	^W 19.91 <u>+</u> 1.94 m.y.	Tmy
MS-253	122 ⁰ 12'40" 44 ⁰ 12'45"	Andesite	^W 8.46 <u>+</u> 0.11 m.y.	Tmv
MS-130	122 ⁰ 06'07" 44 ⁰ 13'10"	Basalt	^W 6.2 <u>+</u> 0.2 m.y.	Tmv
MS-17	122 ⁰ 02'00" 44 ⁰ 12'13"	Andesite	^W 6.2 <u>+</u> 0.2 m.y.	Tmv
A-20	122 ⁰ 02'38" 44 ⁰ 16'39"	Basaltic andesite Ash-flow tuff	^W 5.3 <u>+</u> 0.2 m.y.	Трν
MS-205	122 ⁰ 02'55" 44 ⁰ 09'30"	Basaltic andesite	^W 5.06 <u>+</u> 0.06 m.y.	QTv
MS-208	122 ⁰ 06'30" 44 ⁰ 13'00"	Basaltic andesite	^W 3.88 <u>+</u> 0.06 m.y.	Три
MS-132 MS-110	122 ⁰ 00'50" 44 ⁰ 11'46"	Olivine basalt	^W 2.6+6.2 m.y. W2.1 <u>+</u> 0.1 m.y.	QTv
A-77	122 ⁰ 00'48" 44 ⁰ 17'00"	Basaltic andesite	^w 0.68 <u>+</u> 0.04 m.y.	QTv
U-Cougar	122 ⁰ 14'10" 44 ⁰ 07'46"	Basaltic andesite	^p 16.3 <u>+</u> 1.8 m.y.	Tmvi
U-RI-112	122 ⁰ 16'59" 44 ⁰ 06'30"	Andesite	^p 11.5 <u>+</u> 0.5 m.y.	Tmv
U-RI-85	122 ⁰ 16'10" 44 ⁰ 07'41"	Dacitic ash-flow tuff	^p 13.9 <u>+</u> 0.8 m.y.	Tmv
U-Foley	122 ⁰ 10'29" 44 ⁰ 10'49"	Basalt	^W 2.05 <u>+</u> 0.52 m.y.	QTv
U-Tmw-Top	0 122 ⁰ 11'49" 44 ⁰ 12'11"	Andesite	^w 8.93 <u>+</u> 0.34 m.y.	Tmv
U-Tpb	122 ⁰ 11'31" 44 ⁰ 12'32"	Basaltic andesite	^w 8.39 <u>+</u> 0.36 m.y.	Трν
U-BF-5	122 ⁰ 12'30" 44 ⁰ 08'45"	Dacite	^W 9.31 <u>+</u> 0.44 m.y.	Tmdc

Table 1. Radiometric (K/Ar) ages for selected rocks of the Belknap-Foley area

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* References: MS - McBirney and others, 1974; A - Armstrong and others, 1975; U - Unpublished K/Ar data, University of Utah Research Institute, Stanley Evans and Duncan Foley, analysts. ** w = whole rock date; p = plagioclase date.

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Compo- nent	*T-6 <u>Qbh</u>	J-2 Qtv	J-4 Qtv	J-3 Qtv	S-13 	S-16 Tmv	P-7 <u>Tmv</u> ?
	47 0	47 0	10 20	49.80	53.00	53.0	54.25
SiO ₂	47.0	47.8	48.38		1.10		1.28
Ti0 ₂	-	1.63	2.19	1.68		1.20	
^{A1} 2 ⁰ 3	-	15.42	15.47	15.42	18.30	18.80	16.46
Fe203	-	1.70	1.83	2.20	-7.80	8.20	3.08
Fe0	-	9.54	10.36	9.08	_/.00	0.20	5.92
Mno	-	0.21	0.20	0.21	-	-	0.13
Mg0	-	4.43	5.80	4.17	5.20	5.90	4.46
Ca0	-	10.20	8.21	9.88	8.80	9.30	8.79
Na ₂ 0	-	3.50	3.60	3.90	3.80	3.10	3.46
к ₂ Ō	-	1.30	0.64	1.25	0.40	0.25	0.80
P_05	-	0.02	0.43	0.15	-	-	0.23
H ₂ 0	-	1.36	1.39	0.86	_	-	1.58
	al 47.0	97.11	98.50	98.60	98.40	99.75	100.44
	S-15	S-14	S-9	S-12	S-9	S-10	S-17
	Tmv	_Tmv?	<u>Tmv</u>	Tmv	<u>Tmv</u>	<u>Tov?</u>	<u>Tov?</u>
si0 ₂	54.30	54.70	55.80	57.20	58.00	60.90	61.20
Ti02	1.10	1.10	1.10	0.95	1.00	1.10	1.55
A1203	18.80	17.80	18.50	18.20	16.40	17.40	15.20
F ₂ 0 ₃ Fe0	-7.80	8.0	7.5	7.00	9.6	6.80	8.1
Mn0	-	-	-	-	-	-	-
Mg0	5.20	5.90	4.40	4.20	5.90	2.00	2.50
CaO	8.80	8.50	6.60	8.20	4.90	4.80	4.50
Na ₂ 0	3.00	3.90	4.20	3.50	2.80	4.30	4.30
K20	0.15	0.25	1.70	0.40	0.35	1.20	2.00
P205	-	-	-	-	-	-	-
H ₂ 0							
<u> </u>	-	-	-	-	-	-	-

Table 2. Bulk chemical composition of selected rocks of Belknap-Foley area. (Letters at top of each column indicate sample number and map symbol for stratigraphic unit. All values are in weight percent.)

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*References: P-from Peck, 1964; J-from Jan, 1967; T-from Taylor, 1967; S-from Storch, 1978. - 5 -

(Sutter, 1978). The above units occur principally in the Western Cascades province, in the western part of the map (Plate I).

The eastern part of the study area is completely dominated by compact to diktytaxitic basaltic lavas of Quaternary age (units QTv and Qbh on Plate I). These rocks appear to partially fill in the High Cascade graben described by Allen (1966) and Taylor (1978, 1980). Foley Ridge is a tongue of these Quaternary lavas which filled an east-west-trending canyon cut across the western scarp of the graben about 2.0 m.y. ago (unpublished K/Ar age of 2.05+0.52 m.y. by UURI, Evans and Duncan, analysts).

For this study, the Quaternary lavas were split into two units based on lithology and stratigraphic relationships. Lavas assigned to unit QTv(Pliocene to Pleistocene basalts of the High Cascades) occupy topographic depressions which are clearly related to the present geomorphic setting. Nearly identical rocks located at high elevations of the Western Cascades were assigned to unit Tpv because they occupied topography strongly reversed from the present landscape. K/Ar data, where available, tend to support this division. The Qbh unit (Pleistocene to Holocene basalts of the High Cascades) was identified by extreme freshness of the rock, presence of uneroded tumuli and flow structures on flow tops, obvious control by very youthful drainages, and position above various Quaternary units.

K/Ar dates of samples from units QTv and Qbh are sparse and of relatively low precision and accuracy, but enough are available to provide some age control. Poor precision and accuracy is caused by the relatively low content of K_20 and youthful age of these lavas, which cause very low percentages of radiogenic argon relative to atmospheric argon. The oldest K/Ar date in the area-- 2.6 ± 0.2 m.y. -- is for unit QTv (Laurson and Hammond, 1978); the youngest age is 2.05 ± 0.52 m.y. (UURI date previously cited). A single K/Ar date of

- 6 -

 0.68 ± 0.05 m.y. was obtained from lavas mapped as unit *Qbh* (Laursen and Hammond, 1978).

Structural geology

Faults are concentrated in two major north-south-trending zones along Cougar Reservoir and along Horse Creek-McKenzie River. Both zones have en echelon north- to northwest-trending normal faults with significant dip-slip offsets down to the east. The lineament map (Figure 2) shows additional northeasterly trends along the northern margin of the area.

The Horse Creek-McKenzie River fault zone appears to define the western margin of a major north-south-trending graben which has been partially filled by a shield-like platform of late Pliocene and Quaternary High Cascades basaltic lavas and lesser andesitic ejecta (Allen, 1966; Taylor, 1978, 1980). The youngest dated unit with significant offset on the High Cascades graben margin is a 3.88-m.y.-old basaltic andesite (unit Tpv) on Frissel Point (Sutter, 1978), along the north fork of the McKenzie River. The capping lavas (unit Tpv) west of the north fork of the McKenzie River (Plate I) have been dropped down about 900 m (3,000 ft) to the east along a series of north- to northwest-trending en echelon step faults, but the total structural relief on the graben could be much more than this (Taylor, 1980, personal communication). Only minor offsets appear to affect the Quaternary lavas (Plate I).

The Cougar Reservoir fault zone trend is parallel to the High Cascades graben margin and appears to have a similar sense of movement, with Miocene and Oligocene volcanic rocks appearing to be displaced down toward the east across the zone. Miocene volcanic rocks (unit Imv) are the youngest units with proven offset in this zone.

- 7 -



- 8 -

14

Relation of structures to geothermal systems

The distribution of hot springs in the area is related to the two major north-south-trending fault zones discussed previously. Terwilliger Hot Springs and Cougar Hot Springs are located along the Cougar Reservoir lineament, while Belknap, Foley, and Bigelow Hot Springs are located along the western margin of the High Cascade graben. Hot-spring orifices do not, in general, issue from fault zones but from joints in lavas near the faults.

Three hypotheses might explain the apparent relation of faults to hot springs:

- Faults actively control location of hot springs by serving as conduits for circulation of thermal waters.
- Faults serve as passive controls on the location of hot springs by creating major topographic lows which may fortuitously tap sporadic thermal aquifers.
- Some combination of hypotheses one and two controls the distribution of hot springs.

It is difficult to imagine that fault zones as large as those described here could have no influence on circulation of thermal waters. This is particularly true of the western margin of the High Cascade graben, where the faults are quite young and rocks of different lithology are juxtaposed across the faults. It is also true, however, that hot springs are more likely to issue from topographic lows created by fault-shatter zones, so that hypothesis three above is probably the most logical explanation for control of the hot springs.

- 9 -

GEOPHYSICS

Two geophysical studies were available for evaluation for this report. The first was a regional aeromagnetic study (Figure 3) performed by the Oregon State University Geophysics Group. This study, which is discussed in detail by Couch (1978) and Connard (1980), seems, in general, to show a close correspondence between magnetic maxima and topographic highs in the Belknap-Foley area. This is due to the fact that the Pliocene and Pleistocene units found capping the ridges tend to have a higher proportion of magnetically susceptible lavas than the older, underlying Miocene and Oligocene rocks.

Site-specific interpretations of the aeromagnetic data for the study area are not obvious. However, regional interpretation by Couch (1978) and Connard (1980) indicates a possible fault with east side down that is located in the approximate location of the Western Cascade-High Cascade transition zone fault mapped for this report (see section on geology) and that strikes in approximately the same trend. They also interpret the depth to the Curie point isotherm (temperature below which a material ceases to be paramagnetic; $\sim 600^{\circ}$ C) to be greater on the west side of the fault than on the east side of the fault. This prediction matches well with Blackwell and others (1978), whose thermal model of the Cascades estimates a similar depth to the 600° C isotherm.

The second geophysical study in this report is a regional gravity survey also performed by the Oregon State University Geophysics Group (Couch, 1978; Pitts, 1979). Their survey consists of a complete Bouguer gravity anomaly map (Figure 4) and a residual anomaly map (Figure 5), both of which are discussed in detail by Couch (1978) and Pitts (1979). The main feature of both these maps is the steep gravity gradient coincident with the High Cascades-Western Cascades transition zone and the location of local thermal

- 10 -





CONTOUR INTERVAL 200 FEET

SALEM, OREGON 1960 REVISED 1977

FIGURE 4. COMPLETE BOUGUER ANOMALY MAP OF BELKNAP-FOLEY AREA (From Pitts and Couch, 1978) Contour interval 2.0 mgal

> Estimated uncertainty 1.0 mgal Reduction density 2.67 g/cm³ Transverse Mercator Projection Theoretical gravity: IGF (1930)



1960 REVISED 1977 BELKNAP-FOLEY AREA

(From Pitts, 1979) Contour interval 2.0 mgal

Estimated uncertainty 1.0 mgal Reduction density 2.43 g/cm³ Regional components greater than 895 km removed Transverse Mercator Projection Theoretical gravity: IGF (1930)

14

springs. Pitts (1979) interprets this anomaly to represent either a large graben-bounding fault zone with east side down, an area of shallow silicic intrusives, or a possible combination of both. Detailed geologic mapping and possibly deep drilling are needed to further refine geologic modeling based on the foregoing geophysical studies.

WATER CHEMISTRY

During this study, analyses were compiled of four of the five major thermal springs together with analyses of drill-hole waters in the Belknap-Foley area (Table 3). These data indicate that the thermal waters are generally an alkaline, sodium-chloride-rich carbonate water diagnostic of a hotwater-dominated system at depth with elevated reservoir temperatures (Table 4) calculated by methods presented in Appendix A. Preliminary evaluation of the available data indicates the springs may be placed in two groups. The first are the Bigelow, Belknap, and Foley springs, which show similar amounts of silica (60-110 mg/1), Na:K atomic ratios (54.1-78.5), and calculated minimum reservoir temperatures. Other similarities are seen by comparison of relative amounts of ions such as boron, fluoride, and chloride. The second group are those to the west including the Terwilliger spring cluster, Rider Creek and Walker Creek drill-hole waters (Table 3), and possibly the Cougar Reservoir spring, for which no analyses are available. These springs exhibit lower silica (14-50 mg/1), higher Na:K atomic ratios (80-107), lower amounts of chloride and lithium and higher amounts of sulfate ion, and lower calculated reservoir temperatures (Table 4).

Preliminary data are inconclusive at this point; however, the water group near Cougar Reservoir may represent either a dilute species of the springs to the east or a totally different species. Extensive sampling of thermal spring gases and local cold springs and analyses of all waters for isotopes is needed before a definitive study of the thermal regime can be made.

- 15 -

	Belknap Springs	Belknap Springs	Belknap Springs	Belknap Springs (main)	Belknap Springs (east)
Location	16S/6E/11A	16S/6E/11A	16S/6E/11A	16S/6E/11A	16S/6E/11A
Date sampled	'03	'72	3/76	3/76	3/76
Temp. (^O C)	86.7	71	nt	89.0	66.0
рН	nt	7.62	nt	7.6	7.5
Conductance µmhos∕cm	nt	4300	nt	3900	3720
Alkalinity X _h as mg/l HCO ₃ X _c as mg/l CaCO ₃	nt	nt	nt	14 _c	16 _c
Hardness as mg/1 CaCO ₃	nt	nt	nt	541	544
Total dissolved solids	2506	nt	2550	2491	2377
sio ₂	81	96	110	79.9	70.6
Na	364	690	630	525	490
Κ	69.0	15.0	17.0	16.8	15.2
Ca	455	210	210	208	198
Mg	13	0.2	0.29	0.3	0.4
C1	1343	1300	1550	1195	1036
As	nt	0.35	nt	0.24	0.24
В	nt	6.4	3.6	7.6	7.1
Li	nt	0.95	1.3	1.04	0.95
F	nt	1.2	0.88	1.11	0.98
Fe (total)	nt	0.02	tr	0.1	0.1
A1	nt	nt	0.1	tr	tr
HCO3	nt	17	nt	17	19
P0 ₄	nt	0.21	nt	0.27	0.41
so ₄	168	170	150	105	85
NO ₃	nt	nt	nt	tr	0.08
NH ₃	nt	nt	nt	0.19	0.15

Table 3. Spring and well chemistry of the Belknap-Foley area. All measurements are in mg/l, except for pH or as indicated. nt = not tested; tr = trace.

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	Foley Springs	Bigelow Spring	Terwilliger Springs (upper)	Terwilliger Springs <u>(lower)</u>	Terwilliger Springs (upper)
Location	16S/6E/11A	15S/6E/26Ba	17S/5E/20Bb	17S/5E/20Bb	17S/5E/20Bb
Date sampled	3/76	3/76	'73	3/76	3/76
Temp. (^O C)	80.6	61.0	44.0	38.0	42.0
рН	8.0	7.8	7.7	8.4	8.2
Conductance µmhos/cm	4800	3800	2980	2830	2660
Alkalinity X _h as mg/l HCO ₃ X _c as mg/l CaCO ₃	13 _c	18 _c	nt	15 _c	15 _c
Hardness as mg/1 CACO ₃	1284	459	nt	557	484
Total dissolved solids	3333	2566	nt	1892	1763
SiO ₂	60	69	50	46	47
Na	475	540	392	335	320
К	11	17	6.3	7.3	6.8
Ca	494	188	225	210	196
Mg	0.8	1	0.1	0.2	0.2
C1	1304	1148	788	769	693
As	0.21	0.11	nt	0.1	0.1
В	10.0	6.5	5.1	6.4	6.2
Li	0.96	1.1	0.52	0.7	0.64
F	0.81	1.4	0.8	0.86	0.87
Fe (total)	tr	0.1	tr	0.1	0.1
A1	tr	tr	nt	tr	tr
HCO3	nt	nt	19	nt	nt
PO4	0.06	0.32	nt	0.08	0.08
so ₄	550	102	260	192	185
N 0 3	tr	0.02	nt	0.01	tr
NH ₃	0.15	0.39	nt	0.04	0.12

Table 3. Spring and well chemistry of the Belknap-Foley area--Continued All measurements are in mg/l, except for pH or as indicated. nt = not tested; tr = trace.

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- 17 -

Table 3.Spring and well chemistry of the Belknap-Foley area--Continued. All measurements are in mg/l, except for pH or as indicated. nt = not tested; tr = trace.

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Well or spring	Rider Creek Drill Hole	Walker Creek Drill Hole
Location	175/5E/20 Baa	175/5E/8Acd
Date Sampled	7/80	7/80
Temp. ^O C	18.5	15
Ph	7.77	7.37
Conductance umhos/cm	3400	810
Alkalinity Xh as mg/l HCO ₃	nt	nt
Xc as mg/l CaCO ₃	nt	nt
Hardness as mg/1 CaCO ₃	427.5	68.4
Total dissolved solids mg/l	1962	nt
SiO ₂	14(?)	18(?)
Na	449	149
К	< 2.50	< 2.50
Ca	271	26
Mg	0.500	1
C1	925	135
As	< 0.625	< 0.625
В	7.5	4.2
Li	0.49	< 0.050
F	0.9	0.2
FE (total)	0.12	0.29
A1	< 0.625	< 0.625
HCO3	nt	nt
PO4	nt	nt
so ₄	269	nt
NO ₃	nt	nt
NH ₄	< 0.1	0.6

	Belknap Springs	Belknap Springs	Belknap Springs	Belknap Springs (Main)	Belknap Springs (East)	Terwilliger Springs (Lower)	Terwilliger Springs (Upper)
Flow rate liters/min.	284	300	~250	~250	~250	114	200
Measured temperature ^O C	86.7	71	NT	89	66	38	44
Na:K ^O C	226	87	97	107	104	87	74
Na:K:Ca 1/3 β ^Θ C	202	113	121	125	125	104	95
Na:K:Ca 4/3 β ^O C	110	82	85	83	83	52	48
Na:K:Ca Mg corrected ^O C	183	NC	NC	NC	NC	NC	NC
SiO ₂ conductive	126	135	143	124	119	99	102
SiO ₂ adiabatic ^O C	123	131	137	122	117	100	103
SiO ₂ chalcedony ^O C	98	108	116	97	90	68	72
SiO ₂ opal ^O C	7	15	23	179	164	-17	-14

Table 4. Geothermetric calculations* of minimum reservoir temperatures for selected thermal waters of the Belknap-Foley area

*Methodology for calculations presented in Appendix A. NC = not calculated.

	Terwilliger Springs (Upper)	Bigelow Spring	Foley Spring	Ryder Creek Drill Hole	Walker Creek Drill Hole
Flow rate liters/min.	200	7.6	227	pumped	pumped
Measured temperature ^O C	42	61	80.6	18.5	15
Na:K ^O C	86	104	91	34	76
Na:K:Ca 1/3 β ^O C	103	125	106	61	97
Na:K:Ca 4/3 β ^O C	51	85	52	23	54
Na:K:Ca Mg corrected ^O C	NC	120	NC	NC	NC
SiO ₂ conductive ^O C	99	117	111	NC	NC
SiO ₂ adiabatic ^O C	100	116	110	NC	NC
SiO ₂ chalcedony	69	89	82	NC	NC
SiO ₂ opal ^O C	-16	-1	-6	NC	NC

Table 4. Geothermetric calculations* of minimum reservoir temperatures for selected thermal waters of the Belknap-Foley area -- Continued

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*Methodology for calculations presented in Appendix A. NC = not calculated.

GEOTHERMAL-GRADIENT AND HEAT-FLOW DATA*

The temperature-gradient and heat-flow results for the Belknap-Foley area are as shown in Table 5. Included in the table are the township/range-section and latitude and longitude location of each hole. In addition, the hole name, date of logging used, and collar elevation are included for each hole. The bottom hole temperature, maximum depth, corrected temperature gradient, and, where available, corrected heat flow are printed in blue on Plate I. These values are also listed in the table, as are the depth interval and average thermal conductivity used for calculation of the gradient and heat flow. The values are given in SI units. To transform units, the following conversion factors were used: 1×10^{-6} cal/cm² sec (HFU) = 41.84 mWm⁻², 1×10^{-3} cal/cm sec^OC (TCU) = 0.4184 Wm⁻¹K⁻¹, and 1^{O} C/km = $1 \text{ mKm}^{-1} = 18.2^{O}$ F/100 ft. Corrected gradient and corrected heat flow are values for which the topographic effects have been removed. These are significant for many of the sites studied.

The holes are ranked in terms of the quality of the gradient or heat-flow information: high quality (A), good quality (B), marginal quality (C), data with some problems (D), and data for which no useful temperature gradient or heat flow can be estimated (X). All thermal-conductivity measurements were made on cutting samples. Most of the holes shown on the table were drilled specifically for heat-flow studies, and the data quality is relatively high. In general, holes drilled in the Western Cascade rocks give linear gradients below near-surface effects that may vary in depth from 20 to 100 m. Holes 50-150 m deep in High Cascade rocks, such as 15S/6E-11Dc, are often isothermal because of lateral flow of water in the porous young volcanic rocks.

*By D. D. Blackwell, Southern Methodist University, Dallas, Texas.

Twn/Rng- Section		W Long Deg.Min.	Hole # Date	Collar Elev.	Bottom Temp. (°C)	Depth Interval (m)	Avg. TC Wm ⁻¹ K ⁻¹	# TC	Uncorr. Gradient °C/km	Corr. Gradient °C/km	Corr. HF mWm-2	Q HF	
146/ 6E- 32DC	44-18.29	122- 7,26	HOLF MOH 8/ 1/80	999	19, 12	45:9 154.0	1:46	9,	99:2 1.8	75.6	110	B	
155/ 6E- 11DC	44-16.10	122- 3.25	CR-TBR 7/26/77	716	75.20	.0 52.0						×	
165/ 6E- 2CA	44-12.13	122- 2.97	GR-FP 8/ 5/76	70	14.56	100.0 150.0	1.74 .03	11	84.1 1.4	88.3	154	C	
165/ 4E- 14DBB	44-10.05	122-17.50	BH-32 11/26/75	457	10.76	12.5 45.0	1.80 .33		37.8 .4	35.0	63	D	
165/ 5E- 30AAB	44- 9.31	122-14.62	ST DAM 2 8/ 8/79	389	11:74	25.0 61.0	1.32	1	56.3 1.2	53.0	70	C	
165/ 5e- 30abb	44- 9.29	122-14.88	ST ĐAM 3 8/ 8/79	368	-14.16	15:0 85.0	1:33	4	54.0	51.0	68	D.	
165/ 5E- 30ABC	44- 9:13	122-14.98	ST DAM 1 8/ 8/79	368	12.91	45:0 79.7			50.8 3.6	48.0		С	
165∕ 6E- 27BB	44- 9.06	122- 4.69	ER-HC 9/29/76	573	21.56	30.0 150.0	1.57 .05	12	96.2 .9	70.9	111	B	
175/5E- BACD	44- 6.39	122-13.99	WLKR CRK 7/24/80	585		105.0 155.0	(1.59)		54.1 .7	52.0	83	В	
175/ 5E- 208AA	44- 4.90	122-13.84	RIDR CRK 7/31/80	536	24.77	60.0 154.0	1.64 .04	4	128.5 3.6	97.5	159	В	
175/ 6E- 25AD	44- 3.94	122- 1.37	' MOSQ CRK 8/ 1/80	1005	11.06	115.0 152.0	1.55	З	62.8 1.4	73.8	114	С	
185/ 5E- 11BD	44- 1.12	122- 9.81	. REBL CRK 7/31/80		14.40	55.0 155.0	1.55	Э				×	

Table 5. Geothermal-gradient data, Belknap-Foley area, Oregon

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Only one anomalous value is present, and in general the data fall into two groups: those east of the High Cascade-Western Cascade thermal boundary and those west of the boundary. West of the boundary, heat-flow values generally are below 55 mWm^{-2} , while east of the boundary they are generally above 100 mWm⁻² (Blackwell and others, 1978). Typical gradients are $25-35^{\circ}$ C/km and $60-70^{\circ}$ C/km, respectively. The hole with the highest heat-flow value, at Rider Creek, was drilled within half a mile of Terwilliger Hot Springs, indicating that a slightly larger area is associated with the hot springs than is in evidence from the surface manifestations. Obviously, the value itself is biased by its proximity to the hot springs and cannot be considered a regional value.

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CONCLUSIONS AND RECOMMENDATIONS

During the course of this investigation, two major north-south lineaments were found to have close correlation with the distribution of thermal springs and areas of increased heat-flow. Geological mapping revealed that both lineaments are the result of major north-south-trending fault zones and that these fault zones must, to a certain extent, control the flow of geothermal waters.

The available analyses indicate that the thermal waters may be separated into two compositional groups based on total ionic content, ionic ratios, and calculated reservoir temperatures. These two groups show a one-to-one correlation with the aforementioned fault zones: the hotter springs (i.e., Bigelow, Belknap, and Foley) being associated with the McKenzie-Horse Creek fault zone which forms the western margin of the High Cascade graben; and the cooler springs (Cougar and Terwilliger) being associated with the Cougar Reservoir fault zone which lies west of the High Cascade graben margin. This correlation is also seen in heat-flow measurements, with the higher values associated with the McKenzie-Horse Creek fault zone and the lower numbers associated with the Cougar Reservoir fault zone.

This preliminary data analysis indicates that the McKenzie-Horse Creek fault zone may control a higher temperature geothermal resource than the Cougar Reservoir fault zone. Both zones, however, contain geothermal resources which warrant further study. To accomplish a detailed assessment of the geothermal resources, the following steps are recommended:

- Detailed mapping (scale of 1:24,000 or greater) of the McKenzie-Horse Creek fault zone making use of existing 1:15,000 U. S. Forest Service color aerial imagery -- to identify and evaluate active thermal structures along this zone.
- 2. Detailed spring and well sampling and analyses of both hot and cool waters, including isotopic and gas analyses -- to help evaluate reservoir conditions.

- 24 -

- 3. Closely spaced complete Bouguer and residual gravity anomaly studies along the fault zones -- to further refine the gravity anomalies found during previous regional studies and to tie anomalies to mapped structures.
- 4. Resistivity traverses (either dipole-dipole, roving dipole, or telluric) eastwest and north-south along the fault zones -- to further define geothermal aquifers and to locate areas of thermal upwelling and recharge.
- 5. A program of five to ten 500-ft gradient/stratigraphy holes placed at strategic locations -- to refine the evaluation of the Belknap-Foley heat-flow model.
- 6. Five to six 2,000-ft gradient/stratigraphy holes -- to evaluate thermal anomalies and to directly test geothermal aquifers indicated by resistivity traverses and the shallow heat-flow study.
- Feasibility study -- to determine the best method for drilling the very young, loosely consolidated volcanic rocks within the High Cascade graben.

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APPENDIX A

Formulas used in calculations

Na:K (revised):
t
$${}^{O}C = \frac{1217}{\log (Na/K) + 1.483} - 273.15$$
 (Fournier, 1979)

Na:K:Ca:

$$t^{O}C = \frac{1647}{2.24 + F(T)} - 273.15$$
 (Fournier and Truesdell, 1973),

where F (T) = log (Na/K) + [
$$\beta$$
 log (\sqrt{Ca}/Na)],
 β = 1/3 if t> 100^oC, and 4/3 if t <100^oC,
t^oC = calculated reservoir temperature,
and concentrations are expressed in molality.

Magnesium correction ratio:

 $R = \frac{(\text{milliequivalents Mg})}{(\text{milliequivalents Mg}) + (\text{milliequivalents Ca}) + (\text{milliequivalents K})} \times 100$ If R <5 or >50, no calculation was made. For R between 5-50, $\Delta t_{Mg} = 10.66 - (4.7415) (R) + [(325.87) (\log R)^2] - [(1.032 \times 10^5) (\log R)^2/T] - [(1.968 \times 10^7) (\log R)^2/T^2] + [(1.605 \times 10^7) (\log R)^3/T^2],$

where R = magnesium correction ratio expressed in equivalents,

 ${\vartriangle t}_{M\alpha}$ = the temperature correction that is subtracted from

the Na:K:Ca $1/3 \beta$ calculated temperature,

T = Na:K:Ca 1/3 β calculated temperature in ^OK.

Or $\Delta t_{M_{ff}}$ can be obtained by using the graph compiled by Fournier and Potter (1979).

SiO₂ temperature calculations (Fournier and Rowe, 1966):

SiO ₂ (conductive),	$t^{0}C = \frac{1309}{5.19 + \log (SiO_{2})} - 273.15$
SiO ₂ (adiabatic),	$t^{0}C = \frac{1522}{5.75 + \log (Si0_{2})} - 273.15$
SiO ₂ (chalcedony),	$t^{0}C = \frac{1032}{4.69 + \log (SiO_{2})} - 273.15$
SiO ₂ (opal),	$t^{0}C = \frac{731}{4.52 + \log (SiO_{2})} - 273.15,$

where SiO₂ is expressed in mg/l.

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	HOLE		IZDC MDW			
DEPTH METERS	DATE DEPTH FEET	MEASURED: 9/ TEMPERA DEG C	25/80 ITURE DEG F	geothermai Deg C/KM		FT
00000000000000000000000000000000000000	84950627384051628394061728495062738405162 339550627384051628394061728906396295200162 1111231745074906172890639629962950001020000000000000000000000000000000	6.6.6.7.7.7.7.7.7.8.8.8.8.9.9.9.9.9.9.9.9.9.9	44444444444444444444444444444444444444	00000000000000000000000000000000000000	໑໑໑໑໑ຩຠຆຏ຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺	

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Geothermal-gradient data

	LOCATI	ON: SALEM A 145/ 6E-	MS, DREGON	PAGE	2	
DEPTH METERS	HOLE N DATE M DEPTH FEET	AME: WOLF	MDW 1/25/80	geothermal Deg C/KM I	GRADIENT DEG F/100 F	T
92.0 94.0 98.0 98.0 102.0 1004.0 1102.0 1102.0 1112.0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	901.8 9014.1 9084.1 9084.1 9084.1 9084.1 9084.1 9094.0 909.0 900.0 90	$\begin{array}{c} 12.680\\ 12.860\\ 13.0250\\ 13.430\\ 13.610\\ 13.8980\\ 13.500\\ 13.8980\\ 14.370\\ 14.560\\ 14.550\\ 14.520\\ 15.480\\ 15.680\\ 15.680\\ 15.680\\ 16.390\\ 16.820\\ 16.820\\ 16.820\\ 16.820\\ 16.820\\ 16.820\\ 16.820\\ 16.820\\ 16.820\\ 16.820\\ 16.820\\ 16.820\\ 16.820\\ 17.50\\ 17.50\\ 17.610\\ 17.910\\ 17.910\\ 17.910\\ 18.040\end{array}$	54.82 54.515 55.517 55.515 55.517 55.515 55.517 55.515 55.	00000000000000000000000000000000000000	ຎຨຑຎຨຑຎຎຎຎຎຎຎຎຎຎຎຎຎຎຎຎຎຎຎຎຎຎຎຎຎຎຎຎ	

- 36 -



HOLE NUM	165/ 6E/ 2C. BER: CR-FP			è	
	SURED: 9/29/				ODADIENT
DEPTH	DEPTH	TEMPER		DEG C/KM	LI GRADIENT DEG F/100 FT
METERS	FEET	DEG C	DEGF		
5.0	16.4	7.600	45.68	•0	•0
10.0	32+8	7.270	45+09	-66.0	•3•6
15.0	49.2	7+470	45+45	40+0	5•3
20.0	65.6	7:440	45+39	=6+0	*•3
25•0	82•0	7 • 550	45.59	55•0	1+2
30.0	98+4	7•690	45•84	28•0	1.5
35.0	114.8	7.900	46.22	42.0	2•3
40.0	131+2	7 • 990	46+38	18+0	1.0
45.0	147.6	8+150	46+67	32•0	1+8
50.0	164.0	8+260	46.87	22.0	1.2
55 • 0	180•4	8 • 4 4 0	47•19	36•0	5.0
60.0	196•8	8•690	47•64	50•0	2•7
65.0	213.2	8 • 820	47.88	26.0	1.4
70.0	229.6	9+080	48=34	52.0	2•9
75+0	246+0	9+280	48.70	40.0	5.5
80.0	262.4	9.540	49+17	52.0	2.9
85.0	278.8	9•740	49•53	40.0	5.5
90+0	295+2	10+210	50•38	94+0	5.2
95.0	311.6	10.500	50.90	58+0	3.2
100.0	358•0	10+810	51+46	62.0	3•4
105.0	344 • 4	11+080	51 . 94	54+0	3.0
110.0	360.8	11.450	52+61	74+0	4•1
115.0	377•2	11.810	53.26	72.0	4•0 3•5
120.0	393•6	12 • 130	53.83	64+0	
125.0	410.0	12.630	54.73	100.0	5.5
130+0	426 . 4	13+040	55+47	82+0	4.5
135.0	442+8	13.600	56+48	112.0	6•1
140.0	459.2	14+110.	57.40	102+0	<u> </u>
145•0	475.6	14+500	58•10	78.0	
150•0	492.0	14•810	58+66	62•0	3•4

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LOCATION: SALEM AMS, OPEGON 1657 4E-14DBB HOLE NUMBER: DH-32 DATE MEASUPED: 11726/75

DEPTH	DEFTH	TEMPER	ATURE	GEOTHERM	AL GRADHENT
METERS	FEET	DEG C	DEG F	Deg C/rm	DEG F 100 FT
57.0505050505050505050 11157.0397.03957.0359 11157.0397.03957.0359 1111302030505050505050 11113020505050505050505050 11113020505050505050505050505050505050505050	$\begin{array}{c} 16.4\\ 24.6\\ 92.9\\ 41.2\\ 57.4\\ 65.6\\ 82.0\\ 90.2\\ 98.4\\ 106.9\\ 123.0\\ 131.4\\ 123.0\\ 131.4\\ 123.6\\ 131.4\\ 155.8 \end{array}$	$\begin{array}{c} 9.460\\ 9.120\\ 9.4700\\ 9.4700\\ 9.5540\\ 9.5540\\ 9.5200\\ 9.6200\\ 9.6200\\ 10.1000\\ 10.1800\\ 10.4910\\ 10.4910\\ 10.750\\ 10.750\\ 10.750\end{array}$	49.4765 49.4765 49.9755 49.9755 49.9755 49.97555 49.07555555555555555555555555555555555555	9.0 -136.0 76.0 36.0 32.0 32.0 32.0 32.0 32.0 32.0 32.0 32	0540508888880808884848080 0740811118888884890888484840 140888884848484848 1408888848484848484848484848484848484848

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- 40 -



DEPTH METERS	165/ HOLE NAME: S DATE MEASURED:	PERATURE	Geothermal Gradient Deg C/KM Deg F/100 F	T
5.0 10.0 15.0 25.0 25.0 25.0 25.0 25.0 25.0 25.0 2	16.4 9.050 32.8 9.230 49.2 9.360 65.6 9.510 82.0 9.720 98.4 9.970 114.8 10.240 131.2 10.520 164.0 11.070 180.4 11.340 196.8 11.700 200.1 11.740	48.61 48.85 49.12 49.50 49.95 50.43 50.94 51.44 51.93 52.41 53.06	0.0 0.0 36.0 2.0 26.0 1.4 30.0 1.6 42.0 2.3 50.0 2.7 54.0 3.0 54.0 3.0 54.0 3.0 54.0 3.0 54.0 3.0 54.0 2.2	

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DEPTH DE METERS F	LOCATION: SALEM AI 165/5E- HOLE NAME: ST Di DATE MEASURED: 8 PTH TEMPERI SET DEG C	AM 3 / 8/79	L GRADIENT DEG F/100 FT	
20.0 6 20.0 8 30.0 9 35.0 11 40.0 13 40.0 14 50.0 14 50.0 18 60.0 19 60.0 21 70.0 24 70.0 24	.4 10.990 .8 10.260 .2 10.330 .6 10.290 .6 10.330 .4 10.410 .8 10.550 .2 10.730 .4 10.910 .6 10.910 .6 11.180 .4 11.780 .8 11.780 .9.2 12.070 .6 13.000 .4 13.490 .4 13.980 .4 14.160	$\begin{array}{cccccccccccccccccccccccccccccccccccc$	0.0 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9	

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-44 -



DEPTH METERS	Hole Name Date Measi Depth	165/ 5E-30ABC	GEOTHERM	al gradient Deg F/100 FT	
5.00 1150.00 150.000 150.000 150.000 150.0000000000	$\begin{array}{cccccccccccccccccccccccccccccccccccc$	$\begin{array}{cccccccccccccccccccccccccccccccccccc$	-4.0 28.0 8.0 12.0 20.0 26.0 34.0 52.0 58.0 70.0 48.0 48.0 16.0	0.0254714939286493 	

TEMPERATURE, DEG C



	165/ 6E/27BE BER: CR+HC SURED: 9/29/1				
DEPTH	DEPTH	TEMPEI	RATURE	GEDTHERMAL	GRADIENT
METERS	FEET	DEG C	DEG F	DEG C/KM	DEG F/100 F
5.0	16,4	8+120	46.62	• 0	• 0
10.0	32.8	8.360	47.05	48•0	2•6
15.0	49.2	8+930	48.07	114•0	6•3
20.0	65.6	9.340	48+81	.82•0	4.5
25.0	82•0	9•850	49•73	102.0	5•6
30+0	98+4	10•360	50•65	102.0	5•6
35.0	114.8	10.810	51.46	90.0	4+9
40.0	131+2	11.270	52+29	92•0	5•0
45+0	147•6	11.810	53.26	108•0	5.9
50+0	164.0	12•310	54.16	100.0	5.5
55.0	180.4	12.910	55+24	120.0	5•6 4•8
60•0 (F	196•8 213•2	13•350 14•010	56•03 57•22	38•0 132•0	7•2
65+0 70+0	229+6	14•550	58+19	108•0	5•9
75.0	246•0	14+980	58+96	86+0	4 • 7
80+0	262•4	15+610	60.10	126+0	5•9
85+0	278.8	15.830	60.49	44•0	2•4
90.0	295.2	16•330	61 • 39	100+0	5.5
95.0	311.6	16•640	61.95	62+0	3•4
100+0	328.0	17.070	62.73	86•0	4 • 7
105+0	344 • 4	17.700	63•86	126+0	6•9
110.0	360.8	18.200	64.76	100.0	5•5
115.0	377•2	18•720	65•70	104•0	5•7
120.0	393•6	19•160	66•49	88•0	4 • 8
125.0	410.0	19.630	67.33	94•0	5•2
130.0	426•4	20.050	68.09	84•0	4•6
13340	442•8 459•2	20.420	68•76 69•49	74+0 <u>82+0</u>	4•1 4•5
<u>140+0</u> 145+0	475.6	<u>20•830</u> 21•240	70+23	82+0	4•5
150+0	492•C	21.560	70.81	64•0	3•5



	LOCA HOLE DATE DEPTH DEPTH METERS FEET	Tion: Salem Ams, oregon 175/55- Bacd NAME: WLKR CRK MEASURED: 9/24/80 TEMPERATURE DEG C DEG F	geothermal gradient Deg C/KM deg F/100 Ft	
1 50 1	THE ITHE I 11.0 36.1 13.0 42.6 15.0 49.2 17.0 55.8 19.0 62.3 21.0 68.9 23.0 75.4 25.0 82.0 27.0 88.6 29.0 95.1 31.0 101.7 33.0 108.2 35.0 114.8 37.0 121.4 39.0 127.9 41.0 134.5 43.0 141.0 45.0 147.6 47.0 166.7 51.0 167.3 55.0 180.4 57.0 193.5 61.0 200.1 63.0 206.6 65.0 2113.2 67.0 225.7 83.0 272.2 85.0 226.3 77.0 225.7 83.0 272.2 85.0 228.4 87.0 285.4 89.0 291.9 91.0 298.5	12.410 54.34 12.160 53.69 11.940 53.49 11.940 53.28 11.940 53.28 11.570 52.83 11.570 52.83 11.570 52.63 11.570 52.50 11.370 52.50 11.340 52.50 11.340 52.36 11.320 52.36 11.310 52.36 11.320 52.36 11.390 52.50 11.400 52.57 11.400 52.57 11.450 52.61 11.510 52.72 11.660 53.35 11.990 53.65 12.100 53.78 12.390 54.03 12.390 54.28 12.460 54.28	$\begin{array}{cccccccccccccccccccccccccccccccccccc$	

	DEPTH METERS	FEET	DN: Salem A 175/5E- ME: Wlkr CASURED: 9 TEMPER DEG C	MS, OREGON 8ACD CRK /24/80 ATURE DEG F	Pagi Geothermal Deg C/KM)	e 2 Gradient Deg F/100 FT		
1 5 1	93.0 95.0 995.0 101.0 1005.0 1005.0 1113.0 1113.0 1115.0 1119.0 11119.0 0 11119.0 0 11119.0 0 11119.0 0 11119.0 0 11119.0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	05.6273840516283940.6172884950622738 331184.3840516283940.6172884950622738 333333415546283940.6172884950622738 333333333333333940.61728849528851 3333333333333333340.61728849528851 3333333333333333340.61728849528851 33333333333333333333333333333333333	$\begin{array}{c} 12.540\\ 12.540\\ 12.730\\ 12.8800\\ 12.3950\\ 13.950\\ 13.950\\ 13.950\\ 13.950\\ 13.950\\ 13.950\\ 13.950\\ 13.950\\ 13.950\\ 13.950\\ 13.950\\ 13.950\\ 13.950\\ 13.950\\ 13.950\\ 14.250\\ 14.250\\ 14.830\\ 14.830\\ 14.830\\ 14.830\\ 14.830\\ 14.830\\ 15.1220\\ 15.50$	57514 57514 5155 515 515 515 515 515 515	ଡ଼	ณรมอณจากราชอาตาราชอาตายราชอาตาย ขุดที่สามาราชอาตายราชอาตายราชอาตาย ขุดที่สามาราชอาตายราชอาตายราชอาตาย		•

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	DEPTH METERS	LOCATION: SALEM 175/55 HOLE NAME: RII DATE MEASURED: DEPTH TEMPE FEET DEG C	ams, oregon 208aa Dr Crk 9/23/80 IRATURE DEG F	geothermal Deg C/Km	. GRADIENT DEG F/100 FT	
- 53 -	14.000000000000000000000000000000000000	$\begin{array}{cccccccccccccccccccccccccccccccccccc$	337475985720108394116578084299520429558997726827858 339747598575444721678394116578084299520429558997726827858 33213232323233333333333333555585855558889397246827858 33213232333333333333333555558588893966616488	0.0 -1955.0	໑຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺໊໊໊໊	

	LOCA	TION: SALEM A 175/ 5E-	MS, OREGON 20BAA	PA	GE 2		
DEPTH METERS	HOLE DATE DEPTH FEET	INAME I RIDR	CRK /23/80	geotherma Deg c/km	L GRADIENT DEG F/100 1	T	
$\begin{array}{c} 96.0\\ 978.0\\ 1002.0\\ 1002.0\\ 1008.0\\ 1008.0\\ 1102.0\\ 1114.0\\ 1116.0\\ 1202.0\\ 1114.0\\ 1120.0\\ 1202.0\\ 1202.0\\ 1202.0\\ 1302.0\\ 1302.0\\ 1346.0\\ 0\\ 1422.0\\ 1446.0\\ 0\\ 1446.0\\ 0\\ 152.0\\ 152$	94.94 14.94 14.96 14.96 14.96 14.96 14.96 17.28 19.50 19.73 19.73 19.73 19.73 19.73 19.73 19.73 19.73 19.73 19.73 19.73 19.73 19.73 19.73 19.73 19.73 19.73 19.73 19.73 19.74 19.73 19.74 19.75 19.73 19.75	$\begin{array}{c} 17.350\\ 17.950\\ 18.200\\ 18.500\\ 18.500\\ 19.290\\ 19.290\\ 19.550\\ 19.550\\ 20.350\\ 20.350\\ 20.350\\ 20.350\\ 21.950\\ 22.940\\$	63.791 63.731 64.55.666.77.98 66.77.99 66.168 66.99.94 67.11.04 89.04 60.01 77.72 77.74 7.755 6.667 60 80 80 80 80 77.777 77.7777 77.7777 77.7777777777	$\begin{array}{c} 120.0\\ 155.0\\ 145.0\\ 159.0\\ 1259.0\\ 159.0\\ 159.0\\ 159.0\\ 159.0\\ 159.0\\ 159.0\\ 159.0\\ 159.0\\ 159.0\\ 159.0\\ 155.0\\ 1$	ຩຏໟຩຏໟຏຬໞຬຬຬຏໟຏໟຏຏຬ຺຺຺ ຩຏໟຩຏໟຏຬຬຬຬຬຏໟຏໟຏຬຬ຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺຺		

· 54 -



	DEPTH	LOCATIO HOLE NA DATE ME DEPTH	Y: SALEM AN 175/6E-2 ME: MOSQ ASURED: 9/ TEMPERA	15, OREGON 25AD CRK /24/80 27U85	GEOTHERMAL G		
	METERS	FEET	DEG C	DEG F	DEG C/KM DE	G F/100 FT	
- 56 -	11111222223135599000000000000000000000000000000000	455.687.984.95.0627.7984.051.6287.94.061.7284.95.0627.984.051.6287.94.061.7284.95.0627.7984.051.6287.94.061.7288.49.506.111.121.74.111.145.667.90.063.9.62.99.4.061.7288.49.50.627.788.99.00.111.022.022.022.022.022.022.022.022.	$\begin{array}{c} 6.940\\ 6.920\\ 6.920\\ 6.920\\ 6.920\\ 6.920\\ 6.920\\ 6.950\\ 7.100\\ 7.100\\ 7.100\\ 7.100\\ 7.1100\\ 7.1100\\ 7.1100\\ 7.100\\$	፟፟፟፝፝፝፝ጞቔዀዀዀዀዀዀዀዀዀዀዀዀዀዀዀዀዀዀዀዀዀዀዀዀዀዀዀዀዀዀዀ	0.000000000000000000000000000000000000	0507757888888888881118881888888888888888	

DEPTH METERS	LOCA HOLE DATE DEPTH FEET	1857 5E-: NAME: REBL	CRK /30/80	GEOTHERMA DEG C/KM	AL GRADIENT DEG F/100 FT
10.00 150.00 150.00 150.00 150.00 150.00 150.00 110.00	92.604 911974.604 911974.60 911974.60 911974.60 911996.90 911946.00 911996.90 91100 9106.90 9106.90 9106.90 9106.90 9106.00 9106.00 9106.00 9106.00 9106.00 9106.00 9107.00 9106.00 9107.00 91000.00 9100.00 910000000000	$\begin{array}{c} 9.350\\ 9.260\\ 9.260\\ 9.099\\ 9.099\\ 9.099\\ 9.1220\\ 9.1220\\ 9.1220\\ 9.1220\\ 9.1220\\ 9.1220\\ 9.1220\\ 9.1220\\ 9.1220\\ 9.1220\\ 9.1220\\ 10.3900\\ 10.3900\\ 10.3900\\ 10.3900\\ 10.3900\\ 10.1200\\ 112.0030\\ 14.030\\ 14.030\\ 14.120\\ 14.120\\ 14.120\\ 14.290\\ 14.290\\ 14.290\\ 14.32$	8622429676862862842644268642242686 86224296768628628628628642866428664286642866666666	0.0 -16.0 -16.0 -120.0	000111078886660000098409800499044 0.101107886660000998409800499004

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- 57 -

