

OREGON GEOLOGY

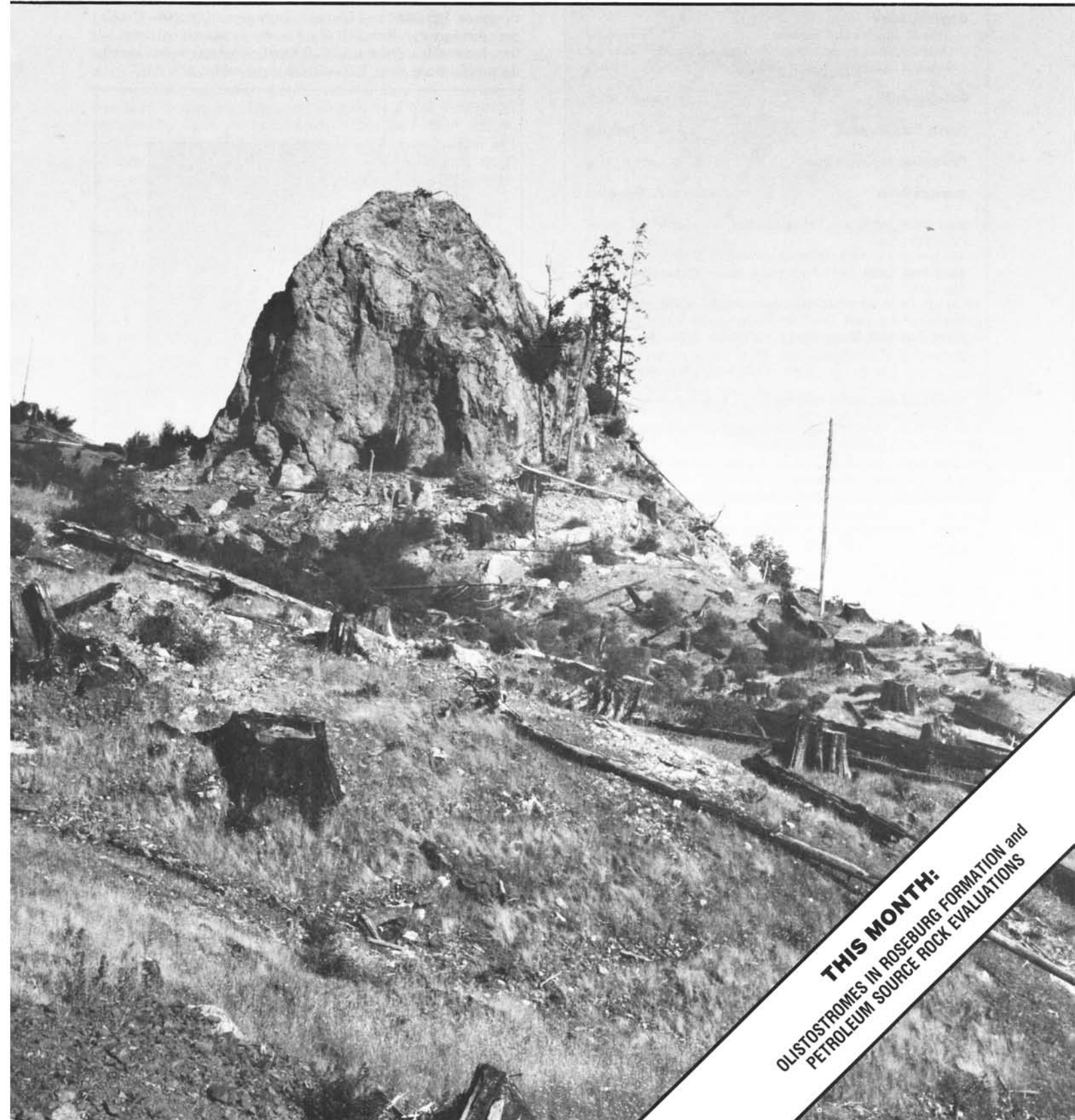
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THIS MONTH:
OLISTOSTROMES IN ROSEBURG FORMATION and
PETROLEUM SOURCE ROCK EVALUATIONS

OREGON GEOLOGY

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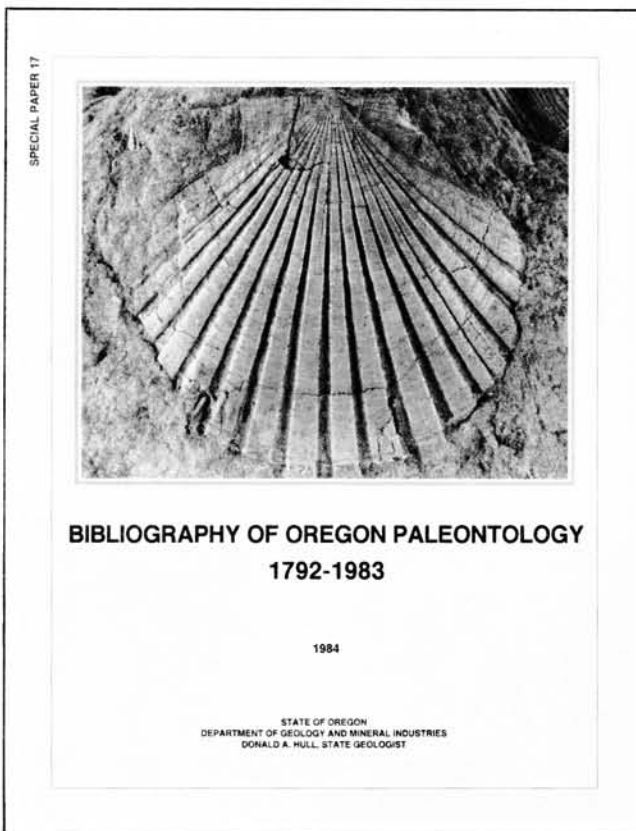
COVER PHOTO

Sandstone pinnacle in Otter Point terrane, upper Sixes River, Curry County, near the Dement Ranch. Article beginning on next page discusses nature of such exotic blocks.

Bibliography of Oregon paleontology released

The Oregon Department of Geology and Mineral Industries (DOGAMI) has released the first comprehensive bibliography of the paleontological literature on Oregon.

Bibliography of Oregon Paleontology, 1792-1983, DOGAMI Special Paper 17, was compiled by Elizabeth L. Orr, Research Librarian in Eugene, Oregon, and William N. Orr, Associate Professor, Department of Geology, University of Oregon. The 82-page publication is the result of six years of research and compilation. It contains approximately 1,200 individual references, including articles from about 160 different serial publications.



Special Paper 17, Bibliography of Oregon Paleontology, by Elizabeth L. and William N. Orr.

The bibliography has two parts: an author index and a subject index. The author entries are annotated, providing information on subject area, location, geologic formation, and similar matters that might not be evident in the title of the entry. The subject index is arranged by formation, county, subject (e.g., Vertebrates), or by a specific name (e.g., Cedar), whenever such a name was used in the title of a reference. Geographic subject terms include broader divisions (e.g., Blue Mountains) as well as smaller subdivisions under the county headings (e.g., Douglas County, Bohemia area). References under individual subject headings also include such geographic subdivisions, so that the literature is accessible through both paleontologic/geologic and geographic subject terms.

The new bibliography, DOGAMI Special Paper 17, is now available at the Oregon Department of Geology and Mineral Industries, 1005 State Office Building, 1400 SW Fifth Avenue, Portland, OR 97201. The purchase price is \$6. Orders under \$50 require prepayment. □

The origin of olistostromes in the Roseburg Formation in southwestern Oregon

by Ewart M. Baldwin, Department of Geology, University of Oregon, Eugene, Oregon 97403

OLISTOSTROMES AND OLISTOLITHS

Olistostromes, as defined in the American Geological Institute (AGI) geologic dictionary (Bates and Jackson, 1980), are sedimentary deposits consisting of a chaotic mass of intimately mixed heterogeneous material (such as blocks and muds) that accumulated as a semifluid body by submarine gravity sliding and slumping of unconsolidated sediments. An olistolith is an exotic block or other rock in an olistostrome. The terms "megabreccia" and "wildflysch" are sometimes used for exotic blocks of unusually large size. A seismoturbidite is an olistostrome produced by an earthquake, and the slide involved would presumably have taken place along a broad front all at one time, in contrast to more random sliding of individual blocks.

There are many exotic blocks of schists, greenstone, chert, basalt, and sandstone within the Roseburg Formation in the Coquille and Sitkum (15-minute) quadrangles east of Myrtle Point and along the Middle, North, and East Forks of the Coquille River near the communities of Gravelford and Bridge.

The writer's attention was first directed to these blocks when he was mapping the Coos Bay 30-minute quadrangle (Allen and Baldwin, 1944). Diller (1901) had located several of these blocks on his geologic folio. Magoon (1966) discovered others in the area he mapped along the edge of the Coquille and Sitkum quadrangles. When compiling the geologic map of Coos County (Baldwin and others, 1973), the writer noted these and other exotic rocks that had been discovered during the extension of many logging roads.

The Otter Point Formation (Koch, 1966; Dott, 1971) of southwestern Oregon is largely a *mélange* of Upper Jurassic and Lower Cretaceous sedimentary rocks, submarine flows, and less abundant pods of chert and blueschists. According to the AGI dictionary, a *mélange* is a mappable body of rock characterized by the inclusion of fragments and blocks of all sizes, both exotic and native, embedded in a fragmental and generally sheared matrix of more tractable material. The definition goes on to speak of a sedimentary *mélange* that is about the same as an olistostrome. The writer uses the term "*mélange*" only to refer to tectonically mixed material. The blocks within the *mélange* are usually more resistant to erosion, so that they stand out as stacks and pinnacles.

Upfaulted blocks of Otter Point terrane are present in the Powers 15-minute quadrangle (Baldwin and Hess, 1971), a short distance south of the Middle Fork of the Coquille River. This terrane contains rocks similar to those now interspersed within the Roseburg Formation and may have provided the source rock for the olistoliths in it.

ORIGIN OF EXOTIC ROCKS

The exotic rocks are common to both the Otter Point and Roseburg terranes, and thus there is a problem in determining which is an Otter Point *mélange* and which is sheared Roseburg along a fault. Such an indeterminate area lies between Powers Junction south of Myrtle Point, Sugarloaf Mountain to the east, and the North Fork of the Coquille River to the north.

The writer mapped a thrust-fault zone extending from the northwest face of Sugarloaf Mountain along the East Fork of the Coquille River to a point just east of the Pleasant Hill School, where it goes under younger and noninvolved Lookingglass sedimentary rocks (Baldwin and others, 1973). Some of the most obvious exotic blocks lie along this zone, and they may be positioned there from thrust Otter Point material or exhumed from nearby Roseburg rock.



Diagram of block movement from the Otter Point terrane to modern sediments.

One large block of greenstone is situated at the south end of the North Fork bridge 2 mi east of Myrtle Point, but it has been largely removed by quarrying. Along the north bank road from this bridge westward to Oregon Highway 42, one soon comes to an outcrop of red chert in the NE $\frac{1}{4}$ sec. 10, T. 29 S., R. 12 W., and, a short distance north of this, an outcrop of yellow sandstone with quartzite-bearing conglomerate which is very close to auger holes that yielded Late Cretaceous microfossils. The other areas between the bridge and the highway are underlain by sheared rock that tends to slide. The writer considered the area to be Roseburg and the exotic rocks to have been introduced by sliding at the time of deposition or by faulting (Baldwin and others, 1973). A greenstone body was present at Kasper's quarry a few miles east of Gravelford, in line with the thrust fault that passes by Sugarloaf Mountain. It is located in the NE $\frac{1}{4}$ sec. 29, T. 28 S., R. 11 W. A few yards down the hill from the quarry are two blocks of schist. Whether both the greenstone and the schist blocks are in Otter Point terrane exposed along the thrust or whether they are olistoliths from nearby Rose-



Block of blueschist near Kasper's quarry along the East Fork of the Coquille River.

burg sediments is difficult to determine. Magoon (1966) mapped the area including Kasper's quarry and suggested the possibilities that the rocks were positioned by faulting or by sliding in, during Roseburg deposition, from a nearby Otter Point highland.

OLISTOSTROME AT BRIDGE, OREGON

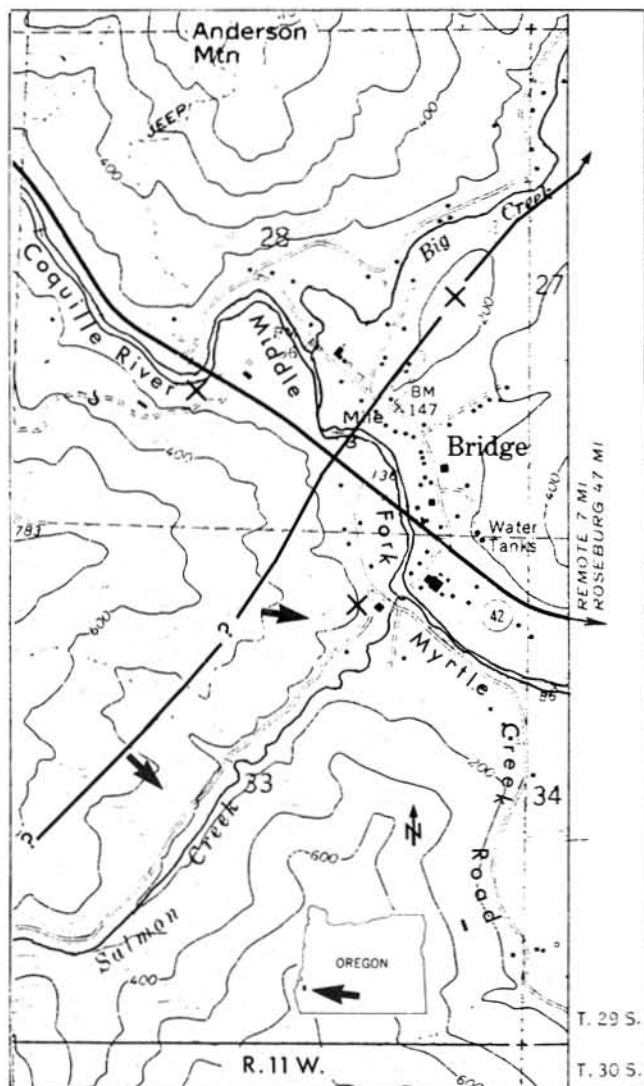
One of the small exotic blocks mapped by Diller (1901) proved to be a block of schist no more than 8 ft in maximum dimension, on a small ridge just north of the business section in the small community of Bridge. Bridge evidently derived its name from the large covered bridge over the Middle Fork of the Coquille River by the school on the Myrtle Creek road. This bridge was not designed for large logging trucks, so when the State Highway Department straightened Highway 42 at this point, creating two bridges over the Middle Fork of the Coquille, the County extended the Myrtle Creek road across the mouth of Salmon Creek to join the highway farther west, thus eliminating the need for the old covered bridge by the schoolhouse. As the Myrtle Creek road was being relocated, most of the material in a cluster of mica-sheet blocks that had been situated on the west bank of Salmon Creek was removed. Another large block of greenstone, 15 ft or more in diameter and almost

metamorphosed to a low-grade blueschist, is situated in the Coquille River just south of the westernmost of the new highway bridges.

The schist block on the small ridge north of the Bridge business section proved to be but one of the exotic rocks in a bed of olistoliths that can be traced to the north across Big Creek and also south to the river bank, then may be projected beneath the alluvial flat and on up the hill southward, paralleling Salmon Creek to the west. There seems to be no doubt as to this bed's being a part of the Roseburg Formation. The bed, which is approximately 4 ft thick, may be seen dipping steeply eastward along the north bank of the river. It contains subrounded to angular blocks of poorly sorted exotic rocks in a sandstone matrix. When projected southward, it goes between the schist blocks at the mouth of Salmon Creek and the large one beneath the highway bridge to the west, but these blocks may have slid from their original positions to where they are now. Other blocks along Salmon Creek probably came from this bed, although some of them may be olistoliths randomly placed in the Roseburg Formation. The writer favors the hypothesis that the olistostrome at Bridge is a seismoturbidite, that is, an olistostrome triggered by an earthquake which caused the olistostrome to form all at one time, in contrast to those singly situated, randomly spaced olistoliths that may have slid into the Roseburg sediments in other parts of the area.

Tor H. Nilsen of the U.S. Geological Survey described the Hilt Bed of the Hornbrook Formation at the Geological Society of America meeting in Indianapolis, November 3, 1983. He traced this bed from Hilt, California, past Ashland, Oregon, and, by projection and intermittent outcrop, to Grave Creek. He considered the Hilt Bed to be a seismoturbidite triggered by an earthquake in Late Cretaceous time. There is a similarity between the Hilt Bed and the one at Bridge, although the latter is not as extensive.

For a modern example of olistostrome formation, we might consider the coastal area south of Humbug Mountain which is underlain by sheared Otter Point terrane (Koch, 1966; Dott, 1971).



Segment of the Bridge 7½-minute quadrangle showing the olistostrome at Bridge and olistolith locations in the vicinity. Black line=olistostrome. X=olistolith location. Arrows=suggested direction of movement of olistoliths.



Sliding terrane with exotic blocks south of Humbug Mountain, Curry County coast.

Here, the unstable sheared material is sliding seaward, with large exotic blocks reaching the shoreline to become stacks and eventually migrating seaward by sliding. Studies of sedimentation in other parts of the world indicate that material may slide for a considerable distance on sea bottoms of only a few degrees slope. As the sea erodes inland, the stacks become situated farther from the shore and surrounded by bedded sediments. In both the modern sediments and the Roseburg Formation the exotic blocks almost surely came from such an unstable Otter Point terrane.

(Continued on page 82, Olistostromes)

Petroleum source rock evaluations of outcrop samples from Oregon and northern California

by B.E. Law, D.E. Anders, T.D. Fouch, M.J. Pawlewicz, M.R. Lickus, and C.M. Molenaar,
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ABSTRACT

Petroleum source rock evaluations of carbonaceous sandstone, siltstone, shale, mudstone, and coal samples of various ages collected from surface exposures in Oregon and northern California indicate that the organic matter is predominantly type III kerogen. Most samples contain less than 1.0 weight percent organic carbon; the average organic carbon content is 0.91 weight percent. The thermal maturities of the samples with respect to hydrocarbon generation range from immature in the younger rocks to post-

mature in the older rocks. Most of the samples are immature to marginally mature. Based on our evaluation of the type of organic matter contained in these samples, indications are the source rock is capable of generating gas and little or no oil.

INTRODUCTION

There has been sporadic interest in the petroleum potential of the Pacific Northwest for several years. Since the 1979 discovery of

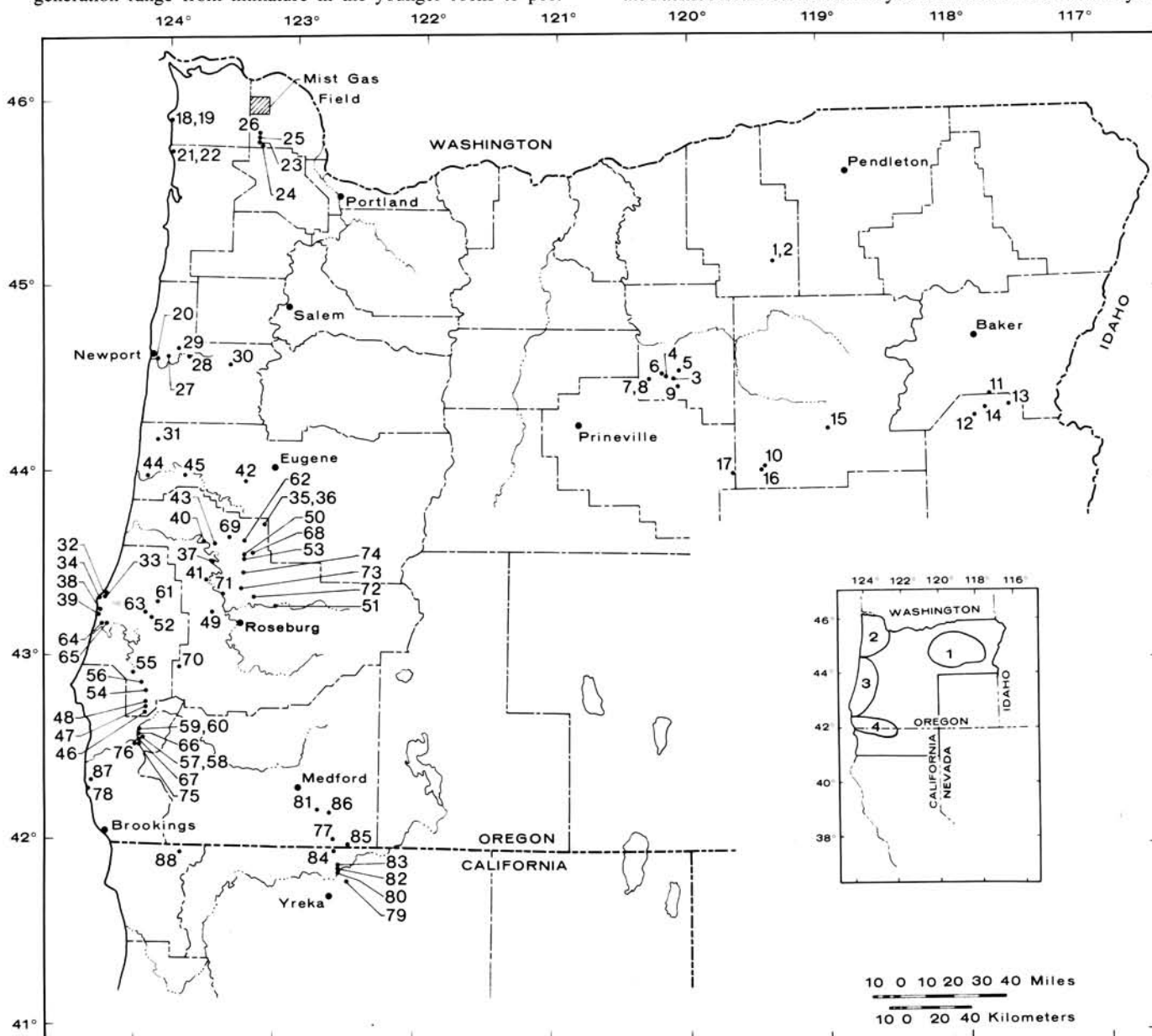


Figure 1. Sample localities and sample areas in Oregon and northern California. Analytical data and more specific locality descriptions for samples are shown in Table 1.

Table 1.--Source-rock analytical data for surface samples from Oregon and northern California

										[Tr. Triassic; JTr. Jurassic and Triassic; J. Jurassic; Cr. Cretaceous; T. Tertiary; argl. argillites; carb. Carbonaceous; clst. claystone; ms. mudstone; ss. sandstone; shv. shale; slst. slate; slstos. siltstone; slty. silt; -- no data]									
Map no.	Sample no.	Section	Township	Range	County	Age	Stratigraphic unit	Lithology	Vitrinite reflectance, R _o (percent)	Organic carbon (wt %)	S ₁ (mg/g)	S ₂ (mg/g)	S ₃ (mg/g)	T _{max} (°C)	Geometric potential (S ₁ +S ₂)	H ₂ index	O ₂ index	Trans-formation ratio, (S ₁ /S ₁ +S ₂)	
Area 1																			
1	0827P-15	SW 34, T. 4 S., R. 28 E., Morrow				K	pre-Clarno Pn.	carb sh	--	4.72	.044	3.83	2.04	434	3.87	81	43	.01	
2	0827P-14	SW 34, T. 4 S., R. 28 E., Morrow				K	Sandstone (Pre-Clarno Pn.)	ss	.86	.04	.017	--	.29	448	.017	--	717	1	
3	0827P-26	6, T. 12 S., R. 22 E., Wheeler				K	Wapinitia Formation	sh	--	1.44	0.013	0.093	0.64	448	0.106	6	44	0.13	
4	0827P-27	35, T. 11 S., R. 21 E., Wheeler				K	Wapinitia Formation	ms	0.62	.70	.028	.12	.40	430	.15	17	58	.19	
5	0827P-28	18, T. 11 S., R. 22 E., Wheeler				K	Wapinitia Formation	ms	.85	.34	.021	.003	.49	449	.024	10	145	.86	
6	0827P-30	28, T. 11 S., R. 21 E., Wheeler				K	Wapinitia Formation	ms	.61	1.19	.017	.123	.76	460	.140	10	64	.12	
7	0827P-31	3, T. 12 S., R. 20 E., Wheeler				K	Wapinitia Formation	ms	.63	.21	.017	.09	.54	438	.697	3	256	.71	
8	0827P-32	3, T. 12 S., R. 20 E., Wheeler				K	Wapinitia Formation	ms	--	.47	.022	.09	.28	426	.11	20	59	.19	
9	0827P-25	21, T. 12 S., R. 22 E., Grant				K	Cable Creek Formation	clst	.84	.84	.014	.006	.53	434	.020	76	63	.69	
10	0827P-22	23, T. 17 S., R. 27 E., Grant				J	Sedimentary rocks	sh	1.06	--	--	--	--	--	--	--	--	--	
11	0827P-16	7, T. 13 S., R. 41 E., Baker				JTr	Sedimentary and volcanic rocks	argl	.49	--	--	--	--	--	--	--	--	--	
12	0827P-17	19, T. 14 S., R. 40 E., Malheur				JTr	Sedimentary and volcanic rocks	argl	--	.99	.017	--	.35	--	.017	--	35	1	
13	0827P-18	32, T. 3 S., R. 42 E., Malheur				JTr	Sedimentary and volcanic rocks	argl	--	1.22	.010	--	.51	--	.010	--	41	1	
14	0827P-19	2, T. 14 S., R. 40 E., Grant				JTr	Sedimentary and volcanic rocks	clst	.50	.22	.009	--	.49	--	.009	--	222	1	
15	0827P-20	11, T. 15 S., R. 31 E., Grant				JTr	Sedimentary and volcanic rocks	sh	3.05	1.00	.016	--	.46	--	.016	--	46	1	
16	0827P-23	18, T. 17 S., R. 27 E., Grant				Tr	Sedimentary and volcanic rocks	slst	.34	--	--	--	--	--	--	--	--	--	
17	0827P-24	25, T. 17 S., R. 27 E., Grant				Tr	Marine sedimentary rocks	sh	.58	--	--	--	--	--	--	--	--	--	
Area 2																			
18	0827P-1	18, T. 5 N., R. 10 W., Clatsop				T	Astoria Formation	clst	0.36	1.18	0.043	0.21	1.33	437	0.25	18	112	0.17	
19	0827P-2	18, T. 5 N., R. 10 W., Clatsop				T	Wapinitia Formation	sh	.65	1.61	.061	.77	.46	424	.83	48	28	.07	
20	0817P-86	SW 9, T. 11 S., R. 11 W., Lincoln				T	Sve Shale (middle part)	slst	.36	1.34	.091	.99	.97	404	1.08	74	72	.86	
21	0827P-3	7, T. 3 N., R. 10 W., Tillamook				T	Formation at Oswald Went	slty clst	1.78	.65	.089	.02	.22	446	.11	2	34	.86	
22	0827P-4	27, T. 4 N., R. 5 W., Columbia				T	Keasay Formation	slst	1.86	--	--	--	--	--	--	--	86	.03	
23	0827P-6	27, T. 4 N., R. 5 W., Columbia				T	Keasay Formation (lower part)	slst	.34	.91	.008	.24	.79	418	.25	27	86	.03	
24	0827P-8	15, T. 4 N., R. 5 W., Columbia				T	Keasay Formation (lower part)	slst, coal	--	.18	.017	.03	.36	401	.05	14	203	.40	
25	0827P-7	3, T. 4 N., R. 5 W., Columbia				T	Coastal Formation (upper part)	slty slst	.31	--	--	--	--	--	--	--	--	--	
26	0827P-7	3, T. 4 N., R. 5 W., Columbia				T	Coastal Formation (lower part)	slty sh	.31	--	--	--	--	--	--	--	--	--	
27	0817P-86	SE NE 5, T. 11 S., R. 10 W., Lincoln				T	Vanilla Formation (lower part)	clst	.32	.50	.021	.17	.51	418	.19	35	103	.11	
28	0817P-84	SW SW 3, T. 11 S., R. 9 W., Lincoln				T	Seelye Formation	sh	.53	.53	.012	.16	.39	422	.17	31	73	.07	
29	0817P-85	SW 24, T. 10 S., R. 10 W., Benton				T	Seelye Formation	sh	.46	.95	.034	.21	.48	427	.24	23	51	.14	
30	0817P-82	SW SW 18, T. 11 S., R. 6 W., Benton				T	Imporia Formation	clst	.53	.23	.012	--	.25	--	.012	--	107	1	
Area 3																			
31	0817P-87	SW NE 10, T. 16 S., R. 11 W., Lane				T	Sve Shale	slst	0.34	1.75	0.046	1.14	0.91	407	1.19	65	52	0.04	
32	0817P-13	NE SW 3, T. 26 S., R. 14 W., Coos				T	Bastardoff Shale (upper part)	slty sh	.32	.94	.087	.99	1.05	418	1.08	106	111	.08	
33	0817P-98	NE SW 3, T. 26 S., R. 14 W., Coos				T	Bastardoff Shale (middle part)	sh	.51	1.46	.017	.22	.51	425	.24	15	35	.07	
34	0817P-4	SE NW 16, T. 21 S., R. 4 W., Douglas				T	Coastal Formation (lower part)	sh	.49	1.44	.016	.39	.76	417	0.41	27	53	.04	
35	0817P-5	SE NW 16, T. 21 S., R. 4 W., Douglas				T	Spencer Formation	clst, carb sh	.48	4.48	.030	.99	1.61	409	1.02	22	36	.03	
36	0817P-3	NE NW 30, T. 23 S., R. 7 W., Coos				T	Wapinitia Formation	sh	.45	1.12	.042	1.09	.24	436	1.06	91	31	.04	
37	0817P-6	C 5, T. 27 S., R. 14 W., Coos				T	Wapinitia Formation	slst, sh	.59	1.35	.026	1.09	.24	436	1.12	81	18	.02	
38	0817P-7	C 5, T. 27 S., R. 14 W., Coos				T	Wapinitia Formation	sh	.45	.86	.029	.37	.66	420	.40	43	47	.07	
39	0817P-8	SW NW 17, T. 27 S., R. 14 W., Coos				T	Wapinitia Formation	sh	.52	.58	.022	.25	.48	424	.27	44	83	.08	
40	0817P-85	SW NW 22, T. 22 S., R. 8 W., Coos				T	Wapinitia Formation	sh	.59	.63	.016	.17	.32	420	.19	27	51	.09	
41	0817P-120	NE SE 36, T. 24 S., R. 8 W., Coos				T	Wapinitia Formation	sh	.59	.91	.018	.51	.33	431	.53	56	36	.03	
42	0817P-91	C NW 29, T. 18 S., R. 5 W., Lane				T	Type Formation (upper part)	slst, sh	.49	.62	.051	--	.39	436	.051	--	95	1	
43	0817P-94	21, T. 22 S., R. 7 W., Douglas				T	Type Formation	slst	.60	1.41	.056	.87	.68	436	.92	62	48	.05	
44	0817P-89	SE SE 19, T. 18 S., R. 11 W., Lane				T	Type Formation	sh	.38	.87	.012	.12	.56	427	.13	14	65	.09	

[JTR, Triassic; JTR, Jurassic and Triassic; J, Jurassic; T, Tertiary; argl, argillite; carb, Carbonaceous; clst, claystone; ms, mudstone; ss, sandstone; sh, sandy sh.; sh, shale; slst, siltstone; slty, silt; -- no data]

Table 1.--Source-rock analytical data for surface samples from Oregon and northern California--continued

Map no. (fig. 1)	Sample no.	Section	Township	Range	County ¹	Age	Stratigraphic unit	Lithology	Vitrinite reflectance, R _o (percent)	Organic carbon (Wt %)	S ₁ (mgHC/g)	S ₂ (mgHC/g)	S ₃ (mgCO ₂ /g)	T _{max} (°C)	Genetic potential (S ₁ +S ₂)	H ₂ index (mgHC/gC)	O ₂ index (mgCO ₂ /gC)	Transformation ratio, (S ₁ /S ₁ +S ₂)
Area 3																		
45	ORITP-90	NE NE 16, T. 18 S., R. 9 W.,	Coos	---	---	T	---	sh	.54	2.55	.042	.96	.93	435	1.00	37	37	.04
46	ORITP-104	NW 21, T. 33 S., R. 11 W.,	Coos	---	---	T	---	sh	.86	.84	.024	.08	.18	435	.10	10	22	.22
47	ORITP-106	SW NE 32, T. 32 S., R. 11 W.,	---	---	---	T	---	coal	.55	54.97	.285	147.71	7.03	423	148.00	269	13	.002
48	ORITP-109	NE 29, T. 32 S., R. 11 W.,	---	---	---	T	---	coal	.63	36.61	.999	59.69	7.70	416	60.69	163	21	.02
49	ORITP-119	SW SW 33, T. 26 S., R. 7 W.,	Douglas	---	---	T	Tyee Formation (lower part)	sh	.58	1.02	.019	.76	.42	435	.78	75	41	.02
50	ORITP-127	SW NW 8, T. 23 S., R. 5 W.,	---	---	---	T	---	sh	.59	.65	.020	.12	.26	430	.14	19	39	.14
51	ORITP-117	C N 1/2 17, T. 26 S., R. 3 W.,	---	---	---	T	Umpqua Formation	clst	.52	.22	.008	.05	.27	432	.06	22	122	.15
52	ORIM-16	C W 1/2 15, T. 27 S., R. 11 W.,	Coos	---	---	T	Umpqua Formation	shly sh	.51	.66	.016	.16	.38	427	.18	24	57	.09
53	ORITP-126	C SW 20, T. 23 S., R. 5 W.,	Douglas	---	---	T	---	shly sh	.61	.31	.019	.02	.75	440	.04	5	243	.54
54	ORITP-102	NW 5, T. 32 S., R. 11 W.,	Coos	---	---	T	---	sh	.58	.66	.025	.18	.23	431	.21	27	35	.12
55	ORITP-100	C NE 34, T. 30 S., R. 12 W.,	---	---	---	T	Umpqua Formation	sh	.46	.89	.019	.11	.47	431	.13	13	52	.14
(-Lookingglass Formation of Baldwin and Beaulieu, 1973)																		
56	ORITP-101	C 19, T. 31 S., R. 11 W.,	---	---	---	T	---	sltst	.61	.61	.017	.13	.33	426	.15	22	54	.12
57	BELO 2082	NE 5, T. 35 S., R. 11 W.,	Curry	---	---	T	Umpqua Formation	shly sh	.75	.77	.016	.09	.23	428	.11	12	30	.14
(-Lookingglass Formation of Ahmad, 1981)																		
58	BELO 2182	SW 5, T. 35 S., R. 11 W.,	---	---	---	T	---	shly sh	.63	.48	.017	.11	.41	440	.13	23	85	.13
59	BELO 2382	SE SW 19, T. 34 S., R. 11 W.,	---	---	---	T	---	shly sh	.63	.35	.015	.01	.34	439	.03	3	96	.60
60	BELO 2482	NE SW 19, T. 34 S., R. 11 W.,	---	---	---	T	---	shly sltst	.64	.41	.029	.152	.31	439	.181	37	75	.16
61	ORITP-116	C 13, T. 26 S., R. 4 W.,	Coos	---	---	T	Umpqua Formation	shly sh	.53	.37	.015	.04	.37	428	.06	12	99	.26
(-Lower part of Lookingglass Formation of Baldwin, 1974)																		
62	ORIM-2	CN 1/2 17, T. 22 S., R. 5 W.,	Douglas	---	---	T	Umpqua Formation	sh, clyst	.65	.32	.016	.12	.38	422	.14	37	120	.12
63	ORIM-15	NW SE 5, T. 27 S., R. 11 W.,	Coos	---	---	T	---	shly sh	.53	1.47	.010	.05	.40	428	.06	3	27	.18
64	ORIM-8	NE 34, T. 27 S., R. 14 W.,	---	---	---	T	Umpqua Formation	shly sh	.52	.37	.024	.094	.28	447	.118	25	75	.20
(-Roseburg Formation of Baldwin and Beaulieu, 1973)																		
65	ORIM-9	NW 35, T. 27 S., R. 14 W.,	---	---	---	T	---	shly sh	.59	.68	.014	.08	.26	424	.09	12	38	.15
66	BELO 2282	SE NW 31, T. 34 S., R. 11 W.,	Curry	---	---	T	Umpqua Fm. (=Roseburg Formation of Ahmad, 1981)	shly sh	.60	.42	.013	.09	.22	441	.10	22	52	.12
67	BELO 2582	NE NE 7, T. 35 S., R. 11 W.,	---	---	---	T	---	carb shly sltst	1.75	.13	.024	---	.23	---	.024	---	57	1
68	ORIM-1	SW 2, T. 23 S., R. 5 W.,	Douglas	---	---	T	Umpqua Formation	shly sh	.60	.43	.009	.007	.46	426	.016	2	107	.56
69	ORITP-92	S 1/2 SW 8, T. 22 S., R. 6 W.,	---	---	---	T	---	sh	.76	.41	.021	.05	.66	422	.07	13	161	.27
70	ORITP-113	NW 16, T. 30 S., R. 9 W.,	---	---	---	T	---	carb sltst	.48	19.20	.037	11.05	16.66	425	11.09	58	87	.003
71	ORITP-121	SE SE 25, T. 25 S., R. 7 W.,	---	---	---	T	---	shly sh	.59	.63	.028	.24	.24	428	.27	38	38	.10
72	ORITP-122	NE SW 36, T. 25 S., R. 5 W.,	---	---	---	T	---	sh	.54	.81	.018	.003	.87	442	.021	33	108	.87
73	ORITP-124	18, T. 25 S., R. 5 W.,	---	---	---	T	---	sh	.56	.34	.011	.06	.23	431	.07	16	68	.17
74	ORITP-125	SW NW 17, T. 24 S., R. 5 W.,	---	---	---	T	---	sh	.64	.84	.015	.14	.25	426	.16	17	30	1
75	BELO 2682	SW NW 18, T. 35 S., R. 11 W.,	Curry	---	---	K	Unnamed Cretaceous rocks	shly sh	.75	2.42	.067	.79	.24	448	.86	33	10	.08
76	BELO 2782	NE 13, T. 35 S., R. 12 W.,	---	---	---	K	---	shly sh	.72	2.48	.040	.36	.62	454	.40	14	25	.10
Area 4																		
77	BELO 1082	NE 32, T. 40 S., R. 2 E.,	Jackson	---	---	T	Colestin Formation (Elliott, 1971)	carb shly ms	0.52	0.11	0.002	---	0.31	---	0.002	---	277	1
78	BELO 282	SW SE 18, T. 38 S., R. 14 W.,	Curry	---	---	K	Hornbrook Formation (Wells and Peck, 1961)	shly sh	.54	.57	.008	.12	.33	431	.13	21	59	.06
79	BELO 482	SE 23, T. 46 N., R. 6 W.,	Siskiyou (CA)	---	---	K	Hornbrook Fm. - Unit b (Nilsen et al., 1983)	coaly sh	.59	23.61	.120	44.68	5.50	442	44.80	189	23	.003
80	BELC 582	NE 32, T. 47 N., R. 6 W.,	---	---	---	K	---	carb sltst, ss	.51	.93	.026	.25	.19	437	.28	27	21	.09
81	BELO 1482	NW 5, T. 39 S., R. 1 E.,	Jackson	---	---	K	---	shly sltst	.77	.64	.012	.04	.33	431	.05	8	65	.21
82	BELC 682	NE NE 29, T. 47 N., R. 6 W.,	Siskiyou (CA)	---	---	K	---	shly sh	.58	.37	.007	.01	.46	455	.02	3	123	.4
83	BELC 782	NW 20, T. 47 N., R. 6 W.,	---	---	---	K	---	shly sh	.67	.74	.022	.22	.34	438	.24	30	46	.09
84	BELC 882	SE NW 25, T. 48 N., R. 7 W.,	---	---	---	K	---	shly sh	.64	.90	.013	.17	.20	434	.18	19	22	.07
85	BELC 982	SE 14, T. 48 N., R. 6 W.,	---	---	---	K	---	shly sh	.60	.69	.016	.20	.26	431	.22	30	37	.07
86	BELO 1282	12, T. 39 S., R. 1 E.,	Jackson	---	---	K	---	shly sh	.52	.64	.009	.01	.83	434	.02	2	129	.39
87	BELO 182	SW SW 32, T. 37 S., R. 14 W.,	Curry	---	---	K	Days Creek Formation (Wells and Peck, 1961)	ms	.55	1.08	.015	.03	.65	433	.05	2	60	.36
88	BELC 382	NW 15, T. 18 N., R. 4 E.,	Del Norte (CA)	---	---	J	Galice Formation	slaty sh	5.33	1.14	0	0	.22	---	0	0	19	0

¹All counties are in Oregon, except as noted.

the Mist Gas Field in northwestern Oregon and encouraging gas shows in wells drilled in central Washington, the region has experienced renewed hydrocarbon exploration activity. However, very little information that might be used to establish the oil and gas source rock potential of the region has been published. The only study in Oregon and northern California is an oil and gas investigation report by Newton (1980) in the Coos Bay, Oregon, area.

The purposes of this study are (1) to report the analytical results of a reconnaissance source rock sampling project in Oregon and northern California and (2) to evaluate the quality, quantity, and thermal maturity of organic matter contained in those samples.

SOURCE ROCK EVALUATION

Eighty-eight samples were collected from outcrops, primarily in road cuts and stream drainages. Sample localities are shown in Figure 1 and are listed in Table 1. Based on the areal distribution of samples, the region was subdivided into four areas (Figure 1). Within each area, the samples are listed in approximate order of increasing age. The age of most samples is Cretaceous or younger. A few pre-Cretaceous rock samples were collected, although rocks older than Cretaceous are commonly metamorphosed. The samples consist of carbonaceous sandstone, siltstone, shale, mudstone, and coal (Table 1) and represent depositional environments ranging from nonmarine to marine. The samples were analyzed for vitrinite reflectance (R_o), total organic carbon (weight percent), and Rock-Eval pyrolysis (Table 1). Vitrinite reflectance and Rock-Eval pyrolysis analyses were conducted by the U.S. Geological Survey (USGS), Denver, Colorado, and organic carbon analyses were conducted by Rinehart Laboratory, Arvada, Colorado.

Rock-Eval pyrolysis includes evaluations of the genetic potential ($S_1 + S_2$), organic matter type (hydrogen index [S_2 /organic carbon] versus oxygen index [S_3 /organic carbon]), and thermal maturity (T^*S_2 max. and transformation ratio [$S_1/S_1 + S_2$]). S_1 represents the quantity of volatile hydrocarbons (HC) expelled from rocks held at 250° C for 5 minutes. S_2 measures the quantity of hydrocarbons (HC) released from the rock upon pyrolysis of the kerogen at 250° to 550° C, programmed at 25° C per minute. S_3 is a measure of the amount of pyrolytic carbon dioxide evolved during the heating interval from 250° to 390° C. Detailed explanations of Rock-Eval pyrolysis are given by Espitalié and others (1977) and Tissot and Welte (1978).

The organic matter richness of the samples is generally low. However, the organic matter content of these surface samples may have been significantly reduced by weathering (Leythaeuser, 1973; Clayton and Swetland, 1978). Most samples contain less than 1.0 weight percent organic carbon (Table 1); average organic carbon content, excluding coal, is 0.91 weight percent. According to Dickey and Hunt (1972), a rock must have a minimum organic carbon content of 0.50 weight percent to be an effective hydrocarbon source rock. Organic matter occurs in most samples as disseminated flakes and plant fragments. The more organic-rich rocks are the coals and carbonaceous shales and mudstones.

Hydrogen and oxygen indices from Rock-Eval pyrolysis are plotted in Figure 2. These results show that the organic matter in the analyzed samples is mainly type III kerogen. The three samples that are in proximity to the merged types I and II evolutionary paths are coal and coaly shale (map numbers 47 and 48 from the Eocene Tyee Formation and 79 from the Cretaceous Hornbrook Formation, Figure 1 and Table 1). Some variation in the hydrogen and oxygen indices may result from different proportions of the organic matter types in the rocks. The hydrogen index is also affected by the organic carbon content of the samples; low organic carbon values commonly give artificially low hydrogen index values due to adsorption of the pyrolysis products by mineral matter in the rock. In general, the more hydrogen-rich kerogens (types I and II) are oil and gas source rocks, and the hydrogen-deficient kerogens (type III) are gas source rocks. Thus, nearly all of the samples from this

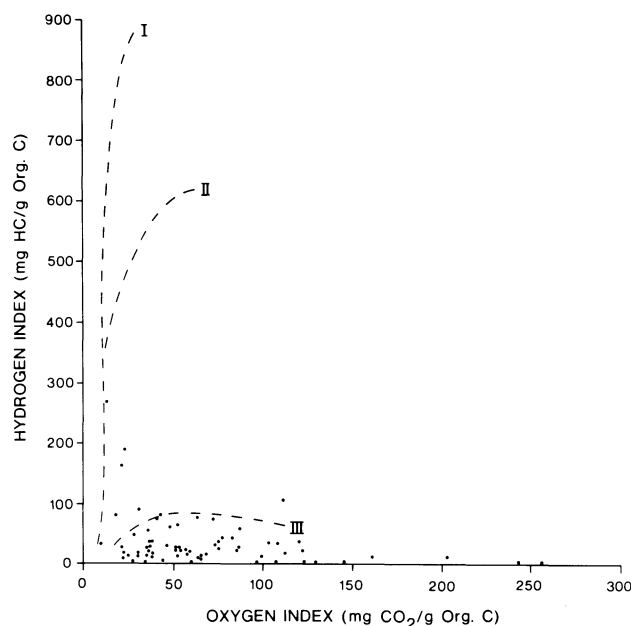


Figure 2. Modified van Krevelen diagram (Tissot and Welte, 1978) of hydrogen and oxygen indices of samples from Oregon and northern California.

study are gas source rocks with little or no potential for oil.

The genetic potential ($S_1 + S_2$ in Table 1) is a measure of the generation capacity of the rock in milligrams of hydrocarbon per gram (mg/g), which is equivalent to kilograms per metric ton (kg/t). Tissot and Welte (1978) have suggested that samples with values less than 2 kg/t have low hydrocarbon potential, samples from 2 to 6 kg/ton have moderate source rock potential, and samples with more than 6 kg/t have good source rock potential. Based on these generation potential values, most of the analyzed samples are poor source rocks, and only a few could be classified as moderate to good source rocks.

The level of thermal maturity at which thermal generation of hydrocarbons starts is variable. The initiation of thermogenic hydrocarbons is largely dependent on the quality of organic matter. According to Tissot and Welte (1978), the top of the "oil window" is between vitrinite reflectance (R_o) values of 0.5 and 0.7 percent, and the bottom of the "oil window" is 1.3 percent. Less is known about the generation of thermogenic gas from type III kerogen, but it is known that significant gas generation begins at a slightly higher level of organic maturation than in types I and II kerogens (Tissot and Welte, 1978; Hunt, 1979). Source rock studies of type III kerogens from Cretaceous and Tertiary rocks in the Greater Green River Basin of Wyoming, Colorado, and Utah (Law and others, 1979, 1980; Law, 1984) indicate that the generation of thermogenic gas begins at a vitrinite reflectance (R_o) value of about 0.80 percent (range is 0.74-0.86 percent). The vitrinite reflectance values of most analyzed samples in this study are below 0.70 percent, indicating that these samples are immature or marginally mature with respect to gas generation.

Temperature maximum (T_{max} in Table 1) data are in fair agreement with vitrinite reflectance data. According to Espitalié and others (1977), temperature maximums in the range of 400° to 435° C correspond to immature source rocks, and temperatures between 435° and 460° C correspond to mature source rocks. The transformation ratio ($S_1/S_1 + S_2$ in Table 1) is another measure of thermal maturity. The transformation ratio generally increases with increasing thermal maturity. Values in the range of 0.1 to 0.4 are considered to be in the oil generation zone (Tissot and Welte,

1978). Coal rank data reported for several coal fields in Oregon by Brownfield (1981) are also in general agreement with the vitrinite reflectance data.

CONCLUSIONS

A source rock evaluation of surface samples collected in Oregon and northern California indicates that the organic matter is predominantly type III kerogen and therefore has the potential of generating gas and little or no oil. The sampling density is insufficient to evaluate the source rock characteristics of any particular stratigraphic unit, but the analyses provide insight into the hydrocarbon potential of the region.

In general, most samples are organically lean, containing less than 1.0 weight percent organic carbon. Thermal maturities of the samples are highly variable. However, most samples are immature to marginally mature with respect to gas generation.

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GSOC meetings announced

The Geological Society of the Oregon Country (GSOC) holds noon luncheon meetings in the Standard Plaza Building, 1100 SW Sixth Ave., Portland, Oregon, in Room A adjacent to the third-floor cafeteria, and 8 p.m. evening lectures at Portland State University, Room 371, Cramer Hall. Upcoming meetings, topics, and speakers are:

July 6 (lecture)—*Rocks, Minerals, and Scenery of Australia*, by Lew and Mura Birdsall, members OAMS and GSOC.

July 13 (lecture)—*Channeled Scablands*, by John Whitmer, M.D., GSOC member.

July 20 (luncheon)—*1984 President's Campout, August 18-26, Lewiston, Idaho*, by Lloyd Wilcox, chairman.

July 27 (lecture)—*Mount St. Helens*, by Bud Beachwood, Public Affairs producer, KOIN-Television.

August 10 (lecture)—*Same topic as July 20: President's Campout*.

For additional information about the lectures or luncheons, contact Viola L. Oberson, GSOC president, phone (503) 282-3685. □

State Capitol mineral display provided by Eugene Mineral Club

The Eugene Mineral Club installed a varied and beautiful display in the State Capitol display case of the Oregon Council of Rock and Mineral Clubs on June 1. Dean and Betty Axtell and Louie and Sibyl Redfern arranged the exhibit provided by 19 members of the club. Lyle Riggs, agent for the Council, and Florence Riggs assisted.

Fourteen Oregon counties are represented by the approximately 70 displayed specimens: Baker, Crook, Grant, Harney, Jackson, Jefferson, Lake, Lane, Lincoln, Linn, Malheur, Sherman, Wasco, and Wheeler Counties.

The display includes three obsidian carvings (a fish, an eagle, and a bear reclining on a small log), petrified wood and limb casts, Paiute agate, natrolite crystal, sagenite, tempskya fern, thunder eggs, Owyhee picture jasper, beach agates, tumbled and faceted sunstones, an agatized clam, carnelian, sunset agate, and a large slab of Chief Paulina agate.

The Eugene Mineral Club display will remain on exhibit until the end of August. □

OIL AND GAS NEWS

Mist Gas Field

Reichhold Energy has extended its exploration activity to the southeast of production by drilling Crown Zellerbach 34-28 in sec. 28, T. 6 N., R. 4 W. The well was spudded April 30 and drilled to a total depth of 3,654 ft. Plugging took place on May 11, 1984.

Columbia County 43-22, completed on February 29, 1984, was put on line May 14, at a rate of approximately 1 MMcf (million cubic feet per day).

Amoco to explore southwest Oregon

On May 25, 1984, the Oregon Department of Geology and Mineral Industries granted three permits to drill to Amoco Production Company (table in June issue) for wells in Douglas and Coos Counties. The wells, two in T. 25 S., R. 9 W. and one in T. 25 S., R. 11 W., are proposed for 15,000 ft and 14,800 ft, respectively. The wells will likely encounter lower to middle Eocene rocks such as the Roseburg and Flournoy Formations. The proposed locations are on land leased from Weyerhaeuser Company.

Columbia County

Champlin Petroleum Company of Denver has been issued a permit to drill in Columbia County (table below). Drilling may start this month (July).

Recent permits

Permit no.	Operator, well, API number	Location	Status, proposed total depth (ft)
263	Champlin Petroleum Co. Puckett 13-36 009-00128	SW ¼ sec. 36 T. 8 N., R. 5 W. Columbia County	Location: 3,500. <input type="checkbox"/>

OSU geology class holds 50th reunion

The first geology class to be graduated from Oregon State University (OSU) met for its fiftieth alumni reunion on June 7, 1984, in Corvallis. As part of the reunion of the OSU class of 1934, the geology graduates met for a social hour at the Dawes House and dinner at the Black Swan Restaurant. Graduates attending included Wayne Felts, Morris Eisenbrey, Leslie Richards, and Robert Gamer. Also attending were geology faculty members including Ira Allison, who was one of the professors for the Class of 1934. ☐

(Olistostromes, continued from page 76)

The Curry County coastline gives us a clue as to how the exotic Otter Point blocks slid into the Roseburg Formation, as the early Eocene seas lapped against the Otter Point highlands.

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Quartzville Corridor saved for recreational mining

Opened officially on May 14, 1984, the Quartzville Recreation Corridor is now reserved for recreational mining and other recreational uses such as camping and fishing. The corridor is a stretch nearly 20 mi along Quartzville Creek in Linn County, beginning at Rocky Top Road on the northeast arm of Green Peter Reservoir and extending upstream to Freezeout Creek Road. It consists mainly of the creek and the land between it and the road which parallels it. Most important for the allowed recreational use of the corridor is the fact that permission for most forms of placering is not needed for the west half of the corridor.

This form of recreational mining in an area that has a rich tradition of mining for gold and silver had been protected in the past by a "mineral withdrawal" provision of the U.S. Bureau of Land Management (BLM) by which the area was kept free of mining claims. As the provision was expiring this year, seven public and private owners and operators of the land involved collaborated to save the corridor for recreation use. The BLM, U.S. Forest Service, Oregon Department of Forestry, City of Sweet Home, and Champion International Corporation, all property owners or managers of the land involved, were brought together with the Linn County Parks Department by the Western Mining Council to work out an arrangement under which Linn County now leases the west half of the corridor and keeps it free of mining claims.



Representatives from several cooperating agencies were on hand for the formal opening of the Quartzville Recreation Corridor. Pictured above, from left to right, are Dave Monson, Public Works Director for the City of Sweet Home; Dave Cooper, Linn County Commissioner; Russell J. Anderson, District Land Manager, Champion International; Jerry Gray (kneeling), resource person for the Western Mining Council; and Merle Marshall, Santiam Area Manager for the Bureau of Land Management. Photo courtesy of The New Era, Sweet Home.

At the entrance to the corridor, a large information sign has been put up by the BLM. Near it, a box contains copies of a brochure printed by Champion International describing the corridor, guidelines for its use, and the mining history of the area.

For additional information about the Quartzville Recreation Corridor, contact the Western Mining Council, Gerald Ullman, 122 Chemawa Road N., Keizer, Oregon 97303, phone (503) 390-3497. ☐

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Available publications

BULLETINS

	Price	No. Copies	Amount
33. Bibliography (1st supplement) geology and mineral resources of Oregon, 1947: Allen	\$ 3.00		
35. Geology of the Dallas and Valsetz quadrangles, rev. 1964: Baldwin (map only)	3.00		
36. Papers on Tertiary foraminifera: Cushman, Stewart, and Stewart, 1949: v. 2	3.00		
44. Bibliography (2nd supplement) geology and mineral resources of Oregon, 1953: Steere	3.00		
46. Ferruginous bauxite deposits, Salem Hills, 1956: Corcoran and Libbey	3.00		
49. Lode mines, Granite mining district, Grant County, Oregon, 1959: Koch	3.00		
53. Bibliography (3rd supplement) geology and mineral resources of Oregon, 1962: Steere and Owen	3.00		
61. Gold and silver in Oregon, 1968: Brooks and Ramp	17.50		
62. Andesite Conference guidebook, 1968: Dole	3.50		
65. Proceedings of the Andesite Conference, 1969: (copies)	10.00		
67. Bibliography (4th supplement) geology and mineral resources of Oregon, 1970: Roberts	3.00		
71. Geology of selected lava tubes in Bend area, Oregon, 1971: Greeley (copies)	5.00		
NEW! 77. Geologic field trips in northern Oregon and southern Washington, 1973	5.00		
78. Bibliography (5th supplement) geology and mineral resources of Oregon, 1973: Roberts	3.00		
81. Environmental geology of Lincoln County, 1973: Schlicker and others	9.00		
82. Geologic hazards of Bull Run Watershed, Multnomah, Clackamas Counties, 1974: Beaulieu	6.50		
83. Eocene stratigraphy of southwestern Oregon, 1974: Baldwin	4.00		
84. Environmental geology of western Linn County, 1974: Beaulieu and others	9.00		
85. Environmental geology of coastal Lane County, 1974: Schlicker and others	9.00		
87. Environmental geology of western Coos and Douglas Counties, 1975	9.00		
88. Geology and mineral resources of upper Chetco River drainage, 1975: Ramp	4.00		
89. Geology and mineral resources of Deschutes County, 1976: Peterson and others	6.50		
90. Land use geology of western Curry County, 1976: Beaulieu	9.00		
91. Geologic hazards of parts of northern Hood River, Wasco, and Sherman Counties, Oregon, 1977: Beaulieu	8.00		
92. Fossils in Oregon (reprinted from <i>The Ore Bin</i>), 1977	4.00		
93. Geology, mineral resources, and rock material of Curry County, Oregon, 1977	7.00		
94. Land use geology of central Jackson County, Oregon, 1977: Beaulieu	9.00		
95. North American ophiolites, 1977	7.00		
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